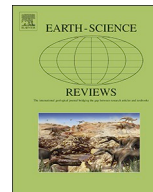




ELSEVIER

Contents lists available at ScienceDirect

Earth-Science Reviews

journal homepage: www.elsevier.com/locate/earscirev

Ancient DNA from marine sediments: Precautions and considerations for seafloor coring, sample handling and data generation



Linda H. Armbricht^{a,*}, Marco J.L. Coolen^b, Franck Lejzerowicz^c, Simon C. George^d, Karita Negandhi^e, Yohey Suzuki^f, Jennifer Young^g, Nicole R. Foster^h, Leanne K. Armandⁱ, Alan Cooper^{a,j}, Martin Ostrowski^{k,l}, Amaranta Focardi^l, Michael Stat^m, John W. Moreau^{n,o}, Laura S. Weyrich^{a,j}

^a Australian Centre for Ancient DNA, School of Biological Sciences, Faculty of Sciences, The University of Adelaide, Adelaide, SA 5005, Australia

^b Western Australian Organic and Isotope Geochemistry Centre (WA-OIGC), School of Earth and Planetary Sciences, Curtin University, Kent Street, Bentley, WA 6102, Australia

^c Center for Microbiome Innovation, University of California, San Diego, La Jolla, CA 92093-0403, USA

^d Department of Earth and Planetary Sciences and MQMarine Research Centre, Macquarie University, North Ryde, NSW 2109, Australia

^e NSW Office of Environment and Heritage, 10 Valentine Avenue, Parramatta, NSW 2150, Australia

^f Department of Earth and Planetary Science, The University of Tokyo, 7-3-1 Hongo, Bunkyo-ku, Tokyo, Japan

^g Advanced DNA, Identification and Forensic Facility, The Environment Institute, Faculty of Sciences, University of Adelaide, Adelaide, SA 5005, Australia

^h School of Biological Sciences, Faculty of Sciences, The University of Adelaide, Adelaide, SA 5005, Australia

ⁱ Australian and New Zealand IODP Consortium, Research School of Earth Sciences, The Australian National University, Acton, ACT 2601, Australia

^j Centre of Excellence for Australian Biodiversity and Heritage, School of Biological Sciences, Faculty of Sciences, The University of Adelaide, Adelaide, SA 5005, Australia

^k Climate Change Cluster, Faculty of Science, University of Technology Sydney, Chippendale, NSW 2007, Australia

^l Department of Molecular Sciences and MQMarine Research Centre, Macquarie University, North Ryde, NSW 2109, Australia

^m School of Environmental and Life Sciences, The University of Newcastle, Callaghan, NSW 2308, Australia

ⁿ School of Earth Sciences, The University of Melbourne, Parkville, VIC 3010, Australia

^o School of Geographical & Earth Sciences, University of Glasgow, Lilybank Gardens, Glasgow G12 8QQ, Scotland

ARTICLE INFO

Keywords:

Ancient DNA
Marine sediments
Deep biosphere
Phytoplankton
Contamination
Seafloor
IODP
Biomarkers
Mars

ABSTRACT

The study of ancient DNA (aDNA) from sediments (*sedaDNA*) offers great potential for paleoclimate interpretation, and has recently been applied as a tool to characterise past marine life and environments from deep ocean sediments over geological timescales. Using *sedaDNA*, palaeo-communities have been detected, including prokaryotes and eukaryotes that do not fossilise, thereby revolutionising the scope of marine micropalaeontological research. However, many studies to date have not reported on the measures taken to prove the authenticity of *sedaDNA*-derived data from which conclusions are drawn. aDNA is highly fragmented and degraded and extremely sensitive to contamination by non-target environmental DNA. Contamination risks are particularly high on research vessels, drilling ships and platforms, where logistics and facilities do not yet allow for sterile sediment coring, and due consideration needs to be given to sample processing and analysis following aDNA guidelines. This review clarifies the use of aDNA terminology, discusses common pitfalls and highlights the urgency behind adopting new standards for marine *sedaDNA* research, with a focus on sampling optimisation to facilitate the incorporation of routine *sedaDNA* research into International Ocean Discovery Program (IODP) operations. Currently available installations aboard drilling ships and platforms are reviewed, improvements suggested, analytical approaches detailed, and the controls and documentation necessary to support the authenticity of aDNA retrieved from deep-sea sediment cores is outlined. Beyond practical considerations, concepts relevant to the study of past marine biodiversity based on *sedaDNA*, and the applicability of the new guidelines to the study of other contamination-susceptible environments (permafrost and outer space) are discussed.

Abbreviations: aDNA, ancient DNA; APC, Advanced Piston Corer; HAPC, Half-Length Advanced Piston Corer; IODP, International Ocean Discovery Program; mbsf, metres below seafloor; MSP, Mission Specific Platforms; NGS, Next generation Sequencing; PCR, polymerase chain reaction; PFT, perfluorocarbon tracer; PMCH, perfluoromethylcyclohexane; PFMD, perfluoromethyldecalin; *sedaDNA*, sedimentary ancient DNA

* Corresponding author.

E-mail address: linda.armbricht@adelaide.edu.au (L.H. Armbricht).

<https://doi.org/10.1016/j.earscirev.2019.102887>

Received 31 October 2018; Received in revised form 31 May 2019

Available online 24 June 2019

0012-8252/ © 2019 Elsevier B.V. All rights reserved.

1. Introduction

Past marine environments have generally been investigated using a suite of methodological approaches and interdisciplinary research fields, such as geology, organic and inorganic geochemistry, paleoceanography and micropaleontology. Discoveries in all of these disciplines have contributed greatly to our understanding of the climatic history of Earth and the evolution and responses of its inhabitants. However, to date, it has not been possible to achieve a detailed picture of all living organisms that have occupied global oceans in the past, restricting estimates of past environmental conditions and climate. The techniques that have traditionally been applied to reconstruct marine palaeo-communities are limited, such as microscopy to investigate the microfossil record (e.g., Winter et al., 2010; Armbrrecht et al., 2018). Due to dissolution and degradation of phytoplankton and microzooplankton while sinking to the seafloor post-mortem, only the most robust skeletons and shells are preserved within a complex geological record (Loucaides et al., 2011). Often, these microfossils are broken, altered by chemical processes and unrecognizable. In the absence of well-preserved diagnostic morphological features, lipid biomarkers can provide supplementary information on biological sources in sediment records (Volkman et al., 1998; Coolen et al., 2004; Sinninghe Damsté et al., 2004; Brocks and Grice, 2011), however, the majority of plankton members do not possess highly diagnostic biomarkers.

New marine metagenomic approaches have allowed the routine characterisation of the diversity of both living hard- and soft-bodied plankton communities in the water column and sub-seafloor. Large-scale “omics” studies, such as the Tara Oceans project (a global sampling program to characterise pro- and eukaryotes of the surface ocean), have shed a new light on our understanding of modern (present day) marine ecosystems and diversity (De Vargas et al., 2015; Sunagawa et al., 2015; Carradec et al., 2018). The deep sea and sub-seafloor have also been targeted with high-resolution metagenomic surveys revealing new insights into the abundance and composition of organisms existing in these largely unexplored environments (e.g., Zinger et al., 2011; Bienhold et al., 2016; Inagaki et al., 2015; Morono and Inagaki, 2016; Orsi et al., 2017, respectively). Such comprehensive studies on living marine communities are continually improving genome reference databases for the hundreds of thousands of pro- and eukaryotic organisms present in the marine environment (Sunagawa et al., 2015; Klemetsen et al., 2017). As a consequence, modern marine metagenomics has not only inspired marine palaeo-research, but also created a means of identifying ancient taxa from marine sediments over geological time-scales.

In the last decade, marine palaeo-research has been reinvigorated by genomic techniques that enable the analysis of ancient DNA (aDNA) molecules from long-dead organisms. Past prokaryotic and eukaryotic plankton communities have been reconstructed using aDNA sequencing approaches (e.g., Coolen and Overmann, 1998, 2007; Coolen et al., 2004, 2008, 2013; Bissett et al., 2005; D'Andrea et al., 2006; Boere et al., 2009; Lejzerowicz et al., 2013; Hou et al., 2014; Randlett et al., 2014; More et al., 2018). These studies have confirmed that phyto- and zooplankton are good targets for aDNA-based studies, while also being particularly relevant for ecosystem-climate reconstructions. It is reasonable to assume that obligate photosynthetic plankton (phytoplankton) and/or zooplankton do not survive and reproduce after burial in deep sediments, and represent uncommon lab contaminants (e.g., Lejzerowicz et al., 2013; Hou et al., 2014; More et al., 2018). aDNA analysis has shown that even after their voyage through the water column plankton-derived particles that had settled on the seafloor still reflect the global biogeographic patterns of living species (Morard et al., 2017). Notably, the reconstruction of past marine communities using aDNA is possible using just a few grams of sediment, facilitating sediment sample collection, transport and storage for the purpose of aDNA analyses.

The marine aDNA archive extends back to the Pleistocene, as shown

by studies of genomic, 18S rRNA gene markers targeting various eukaryotic groups. For example, aDNA has been recovered from various eukaryotic plankton taxa in 43,000-year-old Arabian Sea sediments (More et al., 2018). Taxon-specific approaches targeting small, degraded DNA fragments allowed the retrieval of foraminiferal aDNA from ~800-year-old fjord sediments (Pawłowska et al., 2014) and ~30,000-year-old deep-sea sediments with the additional benefit of enabling the detection of rare taxa (Lejzerowicz et al., 2013). However, if a targeted approach is used, the origin and fate of the DNA in question must be carefully considered, especially for very old claims, such as the retrieval of 1.4 million years old DNA from chloroplasts (Kirkpatrick et al., 2016), which are subject to kleptoplasty (sequestration and maintenance of chloroplasts; Bernhard and Bowser, 1999). While Kirkpatrick et al. (2016) used thorough contamination control, the finding of > 1 million years old DNA remains to be replicated using adapted control measures (e.g., sediment core decontamination and metagenomic sequencing, as outlined in this review). Most studies to date have involved well-dated sediment records and used a cross-validation through paired analysis of aDNA and diagnostic lipid biomarkers as well as geochemical proxies (e.g., Coolen et al., 2006, 2009). Yet, the absence of modern contaminants in analysed samples was not always verified through sequencing analysis of negative sampling and/or extraction controls, which is crucial for the interpretation of aDNA data even if DNA values measured following amplification (by polymerase chain reaction; PCR) are zero (as DNA may be present but simply be below detection limit). To date, the oldest authenticated aDNA records are from ~400,000-year-old cave sediments (Willerslev et al., 2003) and ~700,000-year-old permafrost mammal bones (Orlando et al., 2013).

Despite technologies now being available to rapidly extract and sequence aDNA from marine sediments, and the enormous potential of aDNA research to improve palaeo-oceanographic, –ecosystem and -climate models, marine *sedaDNA* studies remain scarce. This is mainly due to the difficulties and high costs associated with deep-sea aDNA material, for which rarity and hence value justify the deployment of state-of-the-art practices. We review current problems and pitfalls incurred in ship-board sediment sampling, laboratory processing and computational analysis. We suggest solutions to improve sediment coring and sampling strategies so that aDNA research can become a well-established staple in marine biogeosciences. The focus is on sampling protocols within the framework of the International Ocean Discovery Program (IODP) “Biosphere Frontiers” theme, which is dedicated to understanding sub-seafloor communities. Our guidelines for deep-ocean *sedaDNA* isolation are applicable to any low-biomass and setting, including permafrost regions or planet Mars.

2. Definitions and pre-sampling considerations

2.1. Ancient DNA (aDNA), sedimentary ancient DNA (*sedaDNA*), and palaeo-environmental DNA (*PalEnDNA*)

aDNA research involves the biomolecular study of non-modern genetic material preserved in a broad range of biological samples (Shapiro and Hofreiter, 2012; Table 1). When an organism dies, mechanisms that ensure DNA repair in the cell are no longer active, leaving the DNA to degrade over time (Allentoft et al., 2012). Eventually, DNA from dead specimens becomes ancient. aDNA is highly fragmented to an average length of < 100 base pairs (bp), for example, an average length of 48 bp has been determined in the oldest microbial genome assembled to date - from a 48,000-year-old Neandertal (Weyrich et al., 2017). aDNA is affected by post-mortem oxidative and deamination damage, such as thymine enrichment at the end of DNA sequences (Briggs et al., 2007; Ginolhac et al., 2011). Both fragmentation and damage patterns can be used to authenticate aDNA, and damage can even be used to predict its age in certain scenarios (Kistler et al., 2017).

aDNA research mainly focuses on organismal DNA extracted from

Table 1

Terms commonly used in marine *sedaDNA* research and their definition. aDNA terms are listed hierarchically, all other terms are listed alphabetically.

aDNA Terms	Definition	References
Ancient DNA (aDNA)	Biomolecular analysis of non-modern genetic material preserved in a broad range of biological samples.	Shapiro and Hofreiter, 2012
Palaeoenvironmental DNA (PalEnDNA)	Disseminated genetic material found in ancient environmental samples such as sediment, soil, paleosols, coprolites, water and ice.	Rawlence et al., 2014
Environmental DNA (eDNA)	DNA isolated from environmental samples (e.g., soil, water, air), usually a complex mixture of genomic DNA from various organisms.	Taberlet et al., 2012a; Stat et al., 2017
Sedimentary aDNA (<i>sedaDNA</i>)	Ancient DNA isolated from sediments.	Willerslev et al., 2003; Jørgensen et al., 2012
Marine <i>sedaDNA</i>	Ancient DNA retrieved from marine sediment cores.	This review
Other terms used in aDNA research	Definition	References
Background DNA	Unavoidable DNA present in the environment or reagents used to sample and process DNA for sequencing.	Refers to this definition in this review.
Contamination	Biomolecules or living microorganisms introduced from the local environment, sampling procedure, or laboratory reagents into the sample.	Refers to this definition in this review.
Cryosphere	Area of Earth that experiences temperatures below 0 °C for at least part of each year.	Van Everdingen, 1998
Deep Biosphere	Sediment depths below 1 m below seafloor (mbsf).	Parkes et al., 1994; Fry et al., 2008
DNA leaching	Movement of DNA vertically across sediment layers.	Ceccherini et al., 2007; Haile et al., 2007; Pote et al., 2007
Extracellular DNA	DNA released from cells.	Torti et al., 2015
Fossil DNA	DNA from past plankton in the fossil record. Term occasionally used in pioneering studies analysing <i>sedaDNA</i> of plankton and their specific lipid biomarkers in parallel, to corroborate multi-proxy datasets (an analogue would be the term ‘chemical fossils’ (e.g., fossil fuels), occasionally used in organic geochemistry/lipid biomarker/geological studies).	Coolen et al., 2004, 2006, 2009, 2013; Boere et al., 2009, 2011; Coolen and Overmann, 1998, 2007
Intracellular DNA	DNA contained within living cells or structurally intact dead cells.	Torti et al., 2015
Microbiome	All microorganisms present and interacting in an ecosystem.	Blaser et al., 2016
Paleome	Genetic record of past microbial communities acquired from sediments.	Inagaki et al., 2015
Sterile	Contains no living microorganisms (with the debatable exception of prions), but may contain microbial biomolecules, such as DNA or protein.	Refers to this definition in this review.
Tags (also called barcodes or indexes)	Unique DNA sequences that demarcate DNA from a single source.	Refers to this definition in this review.

some tissue remnants of a wide range of single specimen (e.g., tooth, bone, hair, eggshell, feather). In contrast, environmental DNA (eDNA) focuses on disseminated genetic material found in environmental samples such as soil, sediment, water and ice (Taberlet et al., 2012a). Such samples contain complex mixtures of DNA from taxonomically diverse organisms (e.g., bacteria, archaea, plants, animals). In addition to aDNA and eDNA, the term sedimentary aDNA (*sedaDNA*) has been coined to describe aDNA that is exclusively recovered from sediments (Willerslev et al., 2003; Jørgensen et al., 2012). The term fossil DNA has also been used in pioneer studies where sedimentary plankton DNA and lipid biomarkers (i.e., “chemical fossils”) derived from the same historical source organisms were analysed in parallel to validate the ancient DNA results (e.g., Coolen and Overmann, 1998, 2007; Coolen et al., 2004). To a lesser degree, ‘palaeo-environmental DNA’ (PalEnDNA) has also been used to describe disseminated genetic material in a broad range of ancient environmental samples including sediments as well as soil, paleosols, coprolites, water and ice (Rawlence et al., 2014). Modern sequencing technologies and bioinformatic tools ease the analysis of these complex environmental aDNA samples and of the biological responses to human or climate change, with investigations having focussed on terrestrial settings (Jørgensen et al., 2012; Giguët-Covex et al., 2014; Willerslev et al., 2014; Alsos et al., 2015; Pansu et al., 2015). In this review, we use the term ‘marine *sedaDNA*’, which specifically refers to aDNA recovered from ocean sediments. A detailed list of terms frequently used in aDNA research and their definitions is given in Table 1.

2.2. Authenticity of marine *sedaDNA*

2.2.1. Environments favourable for marine *sedaDNA* preservation

Organic-rich sediments deposited in the deep, cold ocean under stratified and anoxic conditions present several favourable

characteristics for the preservation of aDNA (e.g., Coolen and Overmann, 1998, 2007; Coolen et al., 2004, 2013; Boere et al., 2011). Oxidative and deamination damage is reduced in the absence of oxygen (Lindahl, 1993). The absence of irradiation (Lyon et al., 2010), the generally low temperatures (Willerslev et al., 2004), and the high concentration of borate (Furukawa et al., 2013) further contribute to aDNA preservation. Additionally, the typically high mud content of deep-sea sediment offers a particularly well-suited matrix for the preservation and accumulation of DNA (Torti et al., 2015). The high surface:volume ratio of extremely small clay minerals in clay-rich sediments offer a high adsorption surface onto which DNA molecules can bind and remain sheltered from the activity of nucleases (Dell’Anno et al., 2002; Corinaldesi et al., 2008, 2011, 2014, 2018). However, although the above listed properties have been reported to positively impact on DNA preservation, locations with other characteristics that seem less ideal might still be suitable for aDNA research. For example, well-oxygenated Atlantic deep-sea sediments and sand-rich coastal paleo-tsunami deposits have been used to extract and characterise aDNA from foraminifera (Lejzerowicz et al., 2013; Szczuciński et al., 2016, respectively). In conclusion, the preservation of aDNA in marine settings appears to be variable depending on regional environmental characteristics with less favourable to favourable conditions retaining aDNA between a few thousand to, at least, a few ten thousand years. More research is needed to estimate how far back in time authentic marine *sedaDNA* can be detected, which could be achieved, for example, by investigating sediment records from various deep seafloor locations over geological timescales.

2.2.2. Marine *sedaDNA* degradation and fragment length

18S rRNA gene fragments of past dinoflagellates, diatoms, and haptophytes as long as 500 bp in length have been amplified and sequenced (e.g., Coolen et al., 2004), after DNA was isolated from

sediments exhibiting characteristics favourable for aDNA preservation (Section 2.2.1). Up to 20% of genomic DNA from haptophyte algae has been reported to still be of high molecular weight after 2700 years of deposition in Black Sea sediments, and the ratio between 500 bp-long haptophyte 18S rDNA fragments and the concentration of haptophyte-diagnostic long-chain alkenones did not vary substantially for at least 7500 years after deposition, indicative that both types of biomolecules from the same plankton source were equally well preserved (Coolen et al., 2006). This contradicts the generalised view that aDNA is characterised by fragment lengths of < 100 bp. Nevertheless, studies that report the recovery of exceedingly long aDNA fragments should be viewed with scepticism especially in the absence of sampling and extraction controls, where there is no indication on whether the data might reflect modern signals. However, to date, no data are available on average aDNA fragment length for deep-sea sediments, which could be obtained from metagenomic shotgun sequencing. Gaining insights into the latter should be the focus of future research as this information will ultimately help to choose the most suitable and efficient aDNA extraction and sequencing library preparation techniques for degraded *sedadNA* (see Section 3.5).

2.2.3. Contamination sources by modern DNA

Key to the viability of marine *sedadNA* studies is the capability to differentiate between true ancient signals (representative of a particular time-period in the past) and modern contamination (introduced through the sampling and analysis process, or naturally by the environment). Microorganisms and their DNA coat nearly every part of this planet (Weyrich et al., 2015) and a recent study has shown that slow-growing microbes even occur in marine sediments up to 2.5 km deep (Inagaki et al., 2015). The DNA of active deep-biosphere organisms is likely to blur the aDNA signal, as would be the case for microorganisms introduced to ancient sediment samples through the drilling process (see Section 3.2). Moreover, microbial DNA is widely present in laboratory environments and reagents, including in those labelled DNA-free (Salter et al., 2014). If PCR is applied to amplify aDNA, the DNA from modern microorganisms may amplify preferentially over damaged, fragmented aDNA and obscure the true aDNA signals within the sample (Willerslev and Cooper, 2005). Therefore, utmost care must be taken to control and account for contaminants and background DNA throughout the whole process of collecting, processing and sequencing aDNA, e.g., by including negative controls in every step of the analysis process (Fig. 1).

2.2.4. Intracellular vs. extracellular DNA

One approach to separating ancient from modern DNA in sediments has been to differentiate between intracellular and extracellular DNA. Intracellular DNA is defined as DNA contained within living cells, structurally intact dead cells and intact resting stages (e.g., bacterial spores, or other cyst-forming plankton). Extracellular DNA is defined as DNA that has been released from cells and preserved for substantial periods of time through mineral and/or microfossil adsorption or within clay aggregates (Levy-Booth et al., 2007). Extracellular DNA represent an archive of taxa that were autochthonous at the time of deposition (Corinaldesi et al., 2008; Corinaldesi et al., 2011). DNA extraction methods have been developed to target either of these DNA fractions (Corinaldesi et al., 2005; Taberlet et al., 2012b; Alawi et al., 2014). However, it is difficult to prove at what time in the past the organism died, and its DNA became extracellular. Furthermore, the extra- and intracellular DNA pool may not always be clearly distinguishable as genetic material present in the environment might have been taken up by competent bacteria (Demanèche et al., 2001; Dell'Anno et al., 2004) and even by eukaryotes (Overballe-Petersen and Willerslev, 2014). It is also important to note that if only the extracellular pool was to be studied, the paleontological value of dormant yet ancient DNA (e.g., from cysts deposited far back in time) will be lost. Due to these issues, extraction techniques targeting only the

extracellular portion are currently not recommended for marine *sedadNA* studies. Alternatively, bioinformatics approaches that can clearly identify ancient signals (Ginolhac et al., 2011; Kistler et al., 2017) are preferred options for authenticating *sedadNA* sequences (Jónsson et al., 2013).

2.2.5. Vertical DNA movement in marine sediment cores

Three major processes are associated with the vertical movement of DNA in sediment cores: DNA leaching, bioturbation and migration. Bioturbation is a biomechanical process that results in the multi-directional re-organisation of sediments primarily in the upper 10 cm of the sub-seafloor (Boudreau, 1998). DNA leaching is a passive process describing the downward movement of DNA across sediment layers (Haile et al., 2007), without a lowermost boundary. The mixing of sediment layers, and consequently of modern and ancient DNA, can lead to misinterpretations of genomic data. Experimental trials to assess DNA leaching through terrestrial sediments exist (Ceccherini et al., 2007; Poté et al., 2007), with initial results indicating that the extent of leaching depends on the taxonomic source (Haile et al., 2007). Previous studies from lake sediments have shown that leaching is not a factor (Parducci et al., 2017), and in seafloor sediments it seems to play a minor role as aDNA and lipid biomarkers derived from the same microbial source were found to co-exist, or to both be below detection limit, in sediments just centimetres apart (Boere et al., 2009; Coolen et al., 2006, 2009, 2013). In the latter studies indicate that the pore size of the laminated sediments was too small for intracellular DNA to migrate, and that extracellular DNA was adsorbed to the mineral matrices. Recent studies showing *upwards* vertical pore fluid movement also demonstrate the potential for vertical migration of relict or intact DNA within sediments (Torres et al., 2015), and should likewise be considered. Vertical migration of relict or intact DNA is expected to be especially a concern in sediments with micron scale pore sizes and/or a low clay content and a poor capacity to adsorb extracellular DNA. Future experimental research is required to quantify DNA leaching and/or migration through marine sediments, acknowledging the challenge of replicating a complex environmental system varying widely in hydrodynamics and sediment type.

2.2.6. Cross validation of marine *sedadNA* and palaeo-environmental proxies

In addition to using proper contamination controls, downcore changes in past marine community compositions inferred from marine *sedadNA* can be validated through a complementary analysis of independent biological (e.g., microfossils, lipid biomarkers) and geochemical proxies (indicative of the prevailing paleoenvironmental conditions) (Boere et al., 2009; Coolen et al., 2004, 2006, 2013; Hou et al., 2014; More et al., 2018). The most detailed comparison between past ecosystem changes using marine *sedadNA* and the paleo-depositional environment to date has been performed on Holocene sediments from the permanently anoxic and sulfidic Black Sea (Coolen, 2011; Coolen et al., 2006, 2009, 2013; Giosan et al., 2012; Manske et al., 2008). The anoxic and laminated sediments of this semi-isolated sea are devoid of bioturbation and form high-resolution archives of climate-driven hydrological and environmental changes (Calvert et al., 1987; Hay, 1988). Episodes of postglacial sea-level rise ~9000 years ago (Major et al., 2006) and sea surface salinity increase ~5200 years ago (Giosan et al., 2012) have been associated based on *sedadNA* with freshwater to brackish/marine planktonic community transitions (Coolen et al., 2013). For example, the gradual increase in sea surface salinity coincided with the arrival of marine copepods (*Calanus euxinus*), which could only be identified through *sedadNA* analysis (Coolen et al., 2013) as these important zooplankton members generally do not leave other diagnostic remains in the fossil record besides difficult to distinguish resting eggs (Marcus, 1996).

Vice versa, paleoenvironmental conditions inferred from more traditional geochemical and micropaleontological proxies have been

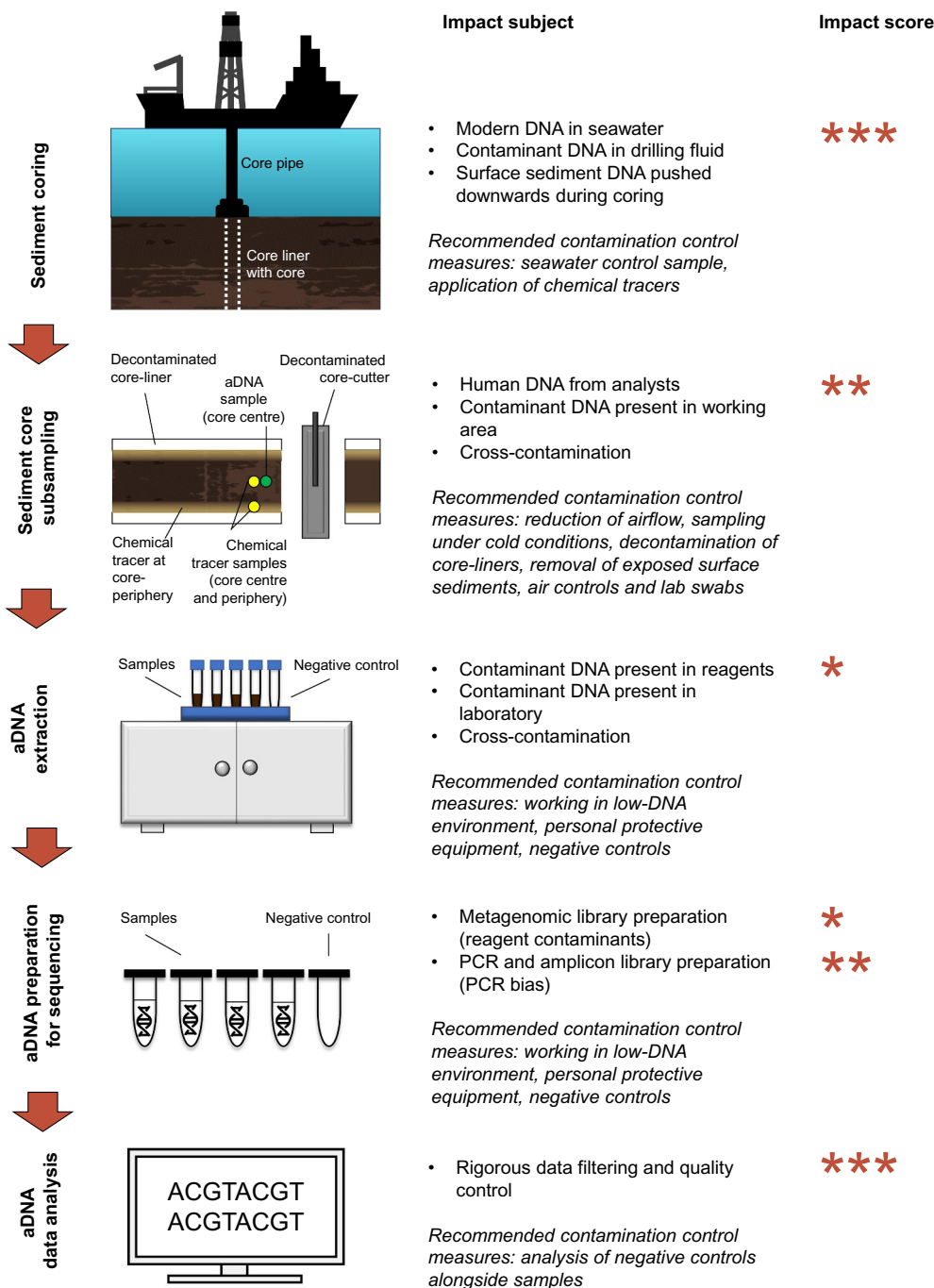


Fig. 1. Schematic showing the key steps involved in acquiring deep marine sediment cores, subsampling, DNA extraction, aDNA preparation for sequencing and data generation. Indicated are sources of potential contamination and reduction in data quality, as well as recommended precautions to be considered and/or controls to be taken. An impact score (1–3 stars) is given to indicate the severity of potential contamination or the impact that impaired data would have on the results at each step in the process. Schematic graphics are not to scale.

verified from parallel *sed*aDNA analysis. By way of example, Black Sea sediments deposited since the last 2500 years contain coccoliths from the calcified marine haptophyte *Emiliania huxleyi* whereas haptophyte-derived diagnostic long chain alkenones in the absence of coccoliths were abundant in up to 7500-year-old sediments (Hay et al., 1991; Coolen et al., 2009). Paired analysis of long-chain alkenones and *sed*aDNA analysis (18S rRNA) revealed that the first haptophytes that colonized the Black Sea ~7500 years ago were initially a mixture of *E. huxleyi* and a highly diverse suite of previously overlooked non-calcified haptophytes related to alkenone-producing brackish *Isochrysis* species. *E. huxleyi* remained the only alkenone producer after 5200 years BP

when salinity reached modern day levels (Coolen et al., 2009). It was concluded that while calcite dissolution prevented the preservation of *E. huxleyi* coccoliths in sediments older than 2500 years ago, their molecular fossils (DNA fragments and long-chain alkenones) survived much longer and showed that in reality this marine haptophyte entered the Black Sea already shortly after the marine reconnection which occurred ~9000 years ago (Coolen et al., 2009, 2013). Even more detailed analyses of *E. huxleyi* (targeting 250-bp-long mitochondrial cytochrome oxidase subunit I; mtCOI) indicate a series of transitions from possibly low-salinity to high-salinity adapted strains of *E. huxleyi* in the Black Sea (7.5–5.2 ka BP), to a different suite of strains during the most

marine stage (5.2–2.5 ka BP), returning to low salinity strains after 2.5 ka BP. The latter transition coincides with the onset of the cold and wet Subatlantic climate (Coolen, 2011) when the Black Sea experienced re-freshening from 32 to 18 ppt (Van der Meer et al., 2008; Giosan et al., 2012; Coolen et al., 2013). The analysis of similar length preserved sequences of viral major capsid protein (mcp) genes revealed a continuous co-existence of *E. huxleyi* and coccolithoviruses in the Black Sea since the last 7000 years and that the same *E. huxleyi* strains, which occurred shortly after the marine reconnection returned with the same viral strains after the re-freshening during the Subatlantic climate thousands of years later (Coolen, 2011). More recently, detailed sedimentary 18S rDNA profiling targeting the shorter (130 bp) V9 region revealed that long-term expansion of past oxygen minimum zones (OMZ) created isolated habitats for unicellular eukaryotes (protists) capable of sustaining oxygen depletion either by adapting a parasitic life cycle (e.g., apicomplexans) or by establishing mutualistic connections with others (e.g., radiolarians and mixotrophic dinoflagellates). These examples show that *seadDNA* can be used to identify biological sources of lipid biomarkers, to verify the reliability of paleoenvironmental information inferred from more traditional proxies, and to reconstruct past ecosystems at multiple trophic levels.

The reconstruction of seafloor prokaryote communities is more complicated since the DNA may be derived from living intact cells in the sediment (see Section 2.2.4). However, 16S rRNA gene profiling from total (intracellular and extracellular) sedimentary DNA has revealed useful insights into sub-seafloor microbial indicators of the palaeo-depositional environment. For example, microbiomes in 20 million years-old coalbeds underlying 2 km of marine sediments were shown to resemble forest soil communities (Inagaki et al., 2015). Variations in bacterial communities found in Baltic Sea sediments have been linked to palaeo-salinity changes (Lyra et al., 2013). Orsi et al. (2017) showed that the genomic potential for denitrification correlated with past proxies for OMZ strength in up to 43 ka-old Arabian Sea sediments. The presence of fermentation pathways and their correlation with the depth distribution of the same denitrifier groups, however, suggests that these microbes were possibly alive upon burial, but low postdepositional selection criteria may explain why they nevertheless formed a long-term genomic archive of past environmental conditions spanning the last glacial-interglacial cycle (Orsi et al., 2017). Further studies are required to determine as to how far the persistence of this phenomenon extends with increased depth in the biosphere. Nevertheless, these examples show that the complementary analysis of marine *seadDNA*-inferred past plankton composition and biological and geochemical proxies is a powerful tool to reconstruct palaeo-environments.

3. *seadDNA* research in the International Ocean Discovery Program (IODP) framework

3.1. IODP infrastructure

IODP is the global community's longest marine geoscience program, operating for 51 years. Its scientific strategy has been to answer globally-significant research questions about the Earth's structure, and the processes that have, and continue to, shape our planet and its climatic history. More recently, additional focus has been cast on biological evolution and limits, particularly in the sub-seafloor environment, under the new Biosphere Frontiers theme (Bickle et al., 2011). This theme has been inspired by the rapidly evolving knowledge and technical capabilities across the multiple merging fields of molecular biology, microbiology, organic and inorganic geochemistry, and micropalaeontology and includes scope for the integration of marine *seadDNA* research. IODP is currently serviced through three platforms, the United States of America's research vessel *JOIDES Resolution*, Japan's *Chikyu* and by the European consortium's Mission Specific Platforms (MSP). In recent years, the laboratories and storage facilities on the ships were modified, or purpose built, to ensure addressing Deep

Biosphere questions was possible. As a result, the latest IODP decadal plan considered options to enable access to uncontaminated samples, their processing and preservation on-board. The latter has led to new coring technologies such as the Half-Length Advanced Piston Corer (HLAPC) allowing a coring depth extension of the conventionally used Advanced Piston Corer (APC), and the use of chemical contamination tracers such as perfluorocarbon tracers (PFTs) (see Sections 3.2 and 3.3, respectively). Particularly useful to aDNA studies may be the development of remotely controlled instruments allowing sediment sampling at ambient pressure (MeBo; Pape et al., 2017) and a rock-drilling device (RD2; Fröh-Green et al., 2015). Notable achievements under the new Deep Biosphere theme include the finding of millions of years old active microbial community from coal beds buried at 2.5 km below the seafloor (Inagaki et al., 2015), and the preservation of an imprint of the Chicxulub impact catastrophe (Cockell et al., 2017). A lot remains to be understood before this theme and its challenges are satisfactorily addressed and it is clear that scientists engaging in Biosphere Frontiers will push methodological, technological and multidisciplinary studies.

3.2. Coring strategies suitable for marine *seadDNA* retrieval

Ideally, marine *seadDNA* sampling involves multiple spatial replicates to ensure that the biodiversity captured is representative of a particular site and time period. However, the ability to collect multiple deep ocean sediment cores to characterise palaeo-plankton is hindered by high costs and logistical issues associated with drilling operations. Thorough planning and collaboration to maximise the use of expensive expeditions and precious deep ocean sediment core material are indispensable in marine *seadDNA* research. To date, several coring strategies exist that differ in machinery as well as sub-seafloor depth that can be reached, and their application is largely dependent on which drilling platform is used (ship or MSP), and what type of sediment is to be cored/drilled (soft sediment or hard rock). This review concentrates on describing piston coring strategies, which are generally better suited to retrieve sediments for aDNA analysis due to relatively low contamination risks. Rotary core barrel systems are required to drill some sedimentary and most igneous rocks, and as they operate with drill-heads and drilling fluids (e.g., seawater) the risk of contamination is dramatically increased (see Section 3.3, Fig. 2).

Piston coring, referred to as Advanced Piston Coring (APC) or Hydraulic Piston Coring System (HPCS), is used to sample unconsolidated or poorly consolidated (i.e., softer) marine sediments. Briefly, these instruments are pushed into the sediment while a piston inside the core pipe creates a vacuum so that the collected sediment remains in the pipe during retrieval. Sediments obtained by piston coring preserve laminated sediments well, are associated with a relatively low risk of environmental contamination and the preferred method to obtain sediments for aDNA analysis (Smith et al., 2000; Lever et al., 2006; More et al., 2018; Fig. 2). Using the piston coring approach, a successive recovery of marine sediments has been achieved to a depth of ~490 m below seafloor (mbsf) (Tada et al., 2013). If only a few metres long (soft) sediment cores are required, gravity-based coring systems, such as a Kasten-, or a Multicorer provide a good alternative (Coolen et al., 2004, 2009). Progress has also been made towards modifying piston coring instruments so that contamination-free sampling is possible, at least for short (< 4 m) sediments (Feek et al., 2011). For example, the 'Mk II sampler' uses an air and water-tight piston coring system with a pointed aluminium head, preventing contamination of the sampled sediment from smearing or water infiltration (Feek et al., 2011). However, to date this corer has only been used in shallow waters, thus it remains to be tested whether use of such an instrument would be feasible during coring operations in deeper waters and which modifications may be required.

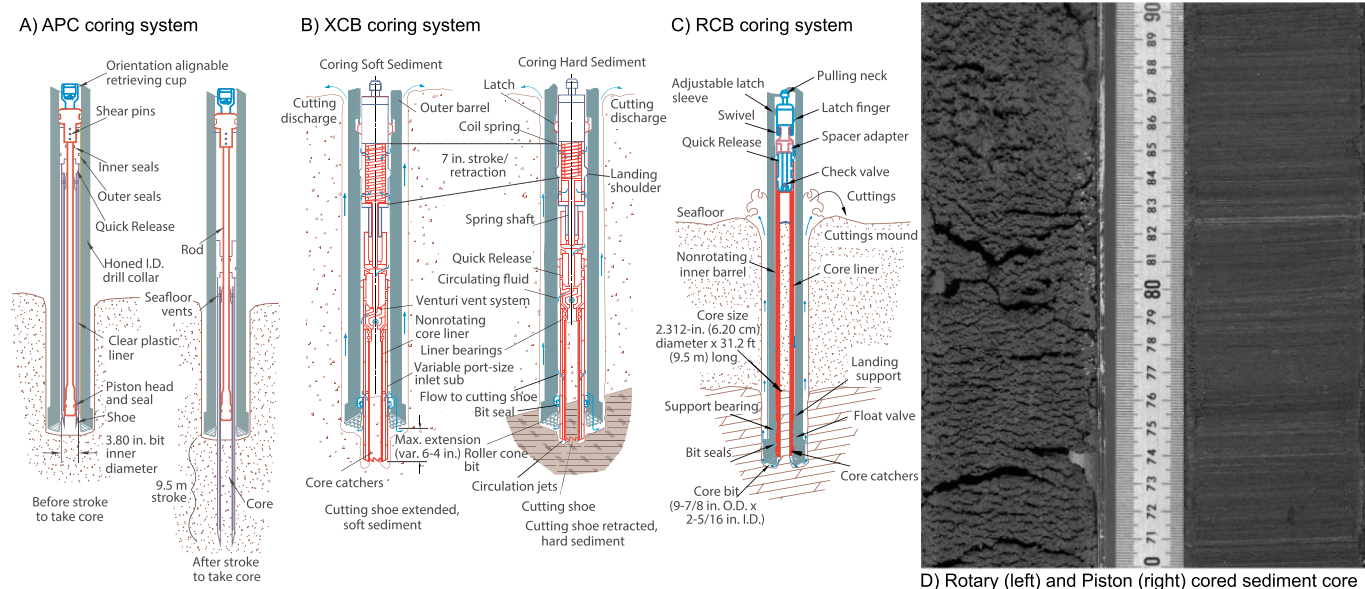


Fig. 2. Overview of IODP coring systems. A) Advanced piston coring system (APC), shown before and after stroking; only small volumes of drill fluid can enter the space between the core barrel and collar from above after stroking, greatly reducing the risk of contamination. B) Extended core barrel system (XCB) and C) Rotary core barrel system (RCB); both containing circulation jets at the bottom of the core barrel through which drill-fluid enters and removes coring debris by transporting it upwards within the drill hole to the surface. D) Comparison of rotary and piston cored sediments demonstrating the well-preserved lamination in Piston cored material. Figure adapted from Sun et al. (2018) and IODP (iodp.tamu.edu/tools/index.html).

3.3. Contamination tracing during coring

Deep ocean coring requires the lowering of coring instruments through hundreds to thousands of metres of seawater before the seafloor is reached, hence exposes the instruments to contamination by modern DNA (Fig. 1). This unavoidable issue has called for the development of methods for environmental DNA contaminant tracing during coring operations. One approach has been to compare biological material found in the contaminating source material (e.g., seawater, drilling fluid) to that of sub-seafloor communities, and to exclude all signals occurring in either from the final analyses (e.g., Expedition 330 Scientists, 2012; Cox et al., 2019). This approach can be implemented for either piston coring or rotary core barrel drilling, provided other sampling constraints associated with these coring systems can be accommodated. However, this procedure does not account for potential “false negative” DNA signals that might indeed occur in both ancient sediments and modern contaminating material. However, in some cases, the microbial community structure of modern contamination (e.g., drilling “mud”) can be resolved, particularly if functional genes are being targeted in sediment samples (Cox et al., 2019).

Another approach has been the introduction of fluorescent microspheres, which are particulate tracers of 0.2–1.0 mm in diameter physically mimicking contaminating organisms. The microspheres have been introduced near the coring head, i.e., where the sediment enters the corer and coring pipe, spreading across the outside of the core (inside the pipe) while drilling, simulating particle movement (Expedition 330 Scientists, 2012; Orcutt et al., 2017). Microscopy has been used to quantify the number of microspheres at the periphery and in the centre of the core to assess contamination (Expedition 330 Scientists, 2012; Orcutt et al., 2017). Similar methods using other perfluorocarbon tracers (PFT's) including perfluoromethylcyclohexane (PMCH) have been developed for the drilling vessel *JOIDES Resolution* (Smith et al., 2000) already in the early phases of IODP. Later, PMCH-based contamination tracing has also been applied during riser drilling on the *Chikyu* (Inagaki et al., 2015). During the IODP Expedition 357 (Atlantis Massif Serpentinization and Life), the PMCH tracer delivery system was further developed to fit the seafloor-based drilling systems MeBO (Pape et al., 2017) and RD2 (Früh-Green et al., 2015) (see

Section 3.1). PMCH is highly volatile which can lead to false positive measurements in uncontaminated samples, therefore, more recent investigations during IODP expeditions have moved to the use of the heavier chemical tracer perfluoromethyldecalin (PFMD, $512.09 \text{ g mol}^{-1}$) (e.g., Fryer et al., 2018).

3.4. Subsampling after core acquisition

Key to enable interdisciplinary sampling and correlations of independent measurements is a detailed sampling plan, specifying sample types as well as the sequence in which these samples are to be collected. Sampling for *sedaDNA* is time-sensitive (to avoid exposure to oxygen, high temperatures and contamination), thus should be conducted immediately after core retrieval on an untreated core-half (i.e., prior to any type of scanning such as by X-Ray). The laboratory in which subsampling for *sedaDNA* is carried out should be clean and workbenches and surfaces decontaminated with bleach (considered to be most efficient at removing contaminating DNA) and, if applicable, ethanol (to prevent corrosion of metal after bleach-treatment). Detailed records on whether molecular and amplification techniques (i.e., PCR) have been employed in on-board laboratories and which organisms were targeted should be kept on record within IODP to ensure sampling for *sedaDNA* can be spatially separated from these laboratories. While most vessels are not currently equipped for complete DNA decontamination, such records may be invaluable for post-expedition *sedaDNA* data analyses.

Two sampling approaches are the most feasible on board IODP ships and MSP's: cutting whole round cores or direct subsampling after core cutting into 1.5 m long sections. The choice of approach needs to be made on a case-by-case basis, and depends on the specific facilities, consumables, chemicals and researcher expertise available during each mission. It is recommended that cutting or subsampling are performed under low or filtered air-flow, e.g., a portable type of a horizontal laminar flow clean air system as described in Morono and Inagaki (2016). Additionally, subsampling should be conducted from the bottom to the top of the core (ancient to modern), using clean (e.g., bleach and ethanol treated) sampling tools for each sample to avoid any form of cross-contamination. Most commonly, soft sediments acquired by piston coring are used for *sedaDNA* analyses, therefore, we focus on

Table 2
Commonly used DNA extraction kits in *sedaDNA* studies to date.

Kit	Reference example	Sediment weight used
DNeasy PowerMax Soil (Mobio/QIAGEN)	Haile, 2012, in: Shapiro and Hofreiter, 2012; Coolen et al., 2013	10 g
DNeasy® PowerSoil® (Mobio/QIAGEN)	Young et al., 2015	0.25 g
DNeasy® PowerLyzer® PowerSoil® (Mobio/QIAGEN)	Leite et al., 2014 (kit comparison)	0.25 g
Fast DNA® Spin Kit (QBIogene)	Haile, 2012, in: Shapiro and Hofreiter, 2012	0.5 g
UltraClean™ Soil DNA Kit Mega Prep (Mobio)	Coolen et al., 2006	0.25 g
Soil Master Kit (Epicentre)	Bidle et al., 2007	pellet

subsampling procedures of the latter here, subsequently briefly outlining sampling recommendations for hard rock material.

If the sampling decision is in favour of whole round core samples, the newly acquired core sections are cut into 5–50 cm sections (preferably under cold conditions), which should be packed in sterile bags or wrap and transferred directly into a fridge or freezer. Although quick and providing a large amount of material for later sub-sampling, this approach has the disadvantage that a lot of freezer space is required, and post-expedition transport can be costly due to the high sample volume and weight.

An alternative to whole round core cutting is direct subsampling immediately after core cutting, either directly from the centre of the top or bottom of each unsplit core section (usually 1.5 m long), or after or splitting the core sections into two halves. In any case the core liner should be cleaned with bleach to remove potential contamination from seawater, and core cutters and splitting-wires, usually metal and sensitive to bleach should be cleaned with ethanol. If sampling from uncut sections, surface material (~0.5 cm) should be removed with bleach and ethanol-treated scrapers before sampling, which is most easily done with sterile cut-tip syringes, placed into sterile plastic bags and stored frozen.

If sampling is undertaken on split core halves, simultaneous visual sedimentological assessments are possible that enable more targeted sampling at specific depths of interest. Using DNA-clean tools, the top 0.5 cm of the core surface should be scraped off perpendicular to the core pipe using sterile scrapers (from bottom to top of the core). Alternatively, the core half to be sampled can be covered with plastic wrap, followed by powdered dry ice, which will result in the top 0.5 cm to become solid frozen. After 5 min, the frozen outer sediment layer can be lifted at one edge with a sterile scalpel creating a contaminant-free surface, from which subsamples can be taken (Coolen and Overmann, 1998). Then, subsampling should be undertaken using sterile (e.g., gamma-irradiated) plastic syringes or centrifuge tubes (e.g., capacity of ~15 mL). Cut-tip syringes have the advantage that more sediment can be collected as no pressure builds up when pushing the syringe into the sediment (the filled syringe should be placed into a sterile plastic bag immediately, e.g., Whirl-Pak®). Alternatively, sterile centrifuge tubes can be used as is to collect 'plunge-samples', usually providing ~1–3 cc of sediment material. The outside of the 'mini-cores' should be cleaned with bleach and placed into sterile plastic bags to avoid cross-contamination between samples. For subsamples, storage at –20 °C or –80 °C is recommended as freezing has been shown to facilitate phytoplankton cell-lysis during DNA extractions (Armbrecht et al., in prep.). Sub-samples can also be collected by transferring a small amount of sediment into a sterile microcentrifuge tube using clean metal or disposable spatulas (particular care needs to be taken to avoid cross-contamination when using the same sampling tool for different samples). The latter approach may be a good solution when only a few small samples are required, e.g., to supplement other scientific questions of an ongoing expedition. For replication purposes it is recommended that duplicate samples are taken at each depth.

If the material is hard rock or similar, subsamples are most easily collected from whole round or split cores. The same decontamination procedures as outlined above should be considered throughout the subsampling procedure (i.e., decontamination of work-surfaces and

sampling tools with bleach and ethanol, sampling under cold conditions and filtered or low air-flow, packing of samples into sterile bags before storage). A de-contaminated metal cutter or a hammer and chisel are best used to remove the outer layer of the exposed sediment, at least at those depths where subsampling is anticipated.

3.5. Marine *sedaDNA* sample processing and analysis

Marine *sedaDNA* samples should be processed in a specialised aDNA laboratory to prevent contamination with modern DNA. Such a laboratory is generally characterised by creating a low-DNA environment, with a clear separation of no-DNA (e.g., buffer preparation) and DNA-containing (e.g., DNA extraction) workflows, regular and thorough sterilisation procedures, positive air pressure, and protective clothing of the analyst (lab coat/suit, gloves, facemask, visor). Details on optimised laboratory set-up, techniques and workflows have been reviewed before (Cooper and Poinar, 2000; Pedersen et al., 2015). The introduction of *sedaDNA* samples into such facilities is relatively straight-forward, as the outer packaging and surface of the sample can be easily sterilised (e.g., using bleach and/or UV).

As on-board subsampling, DNA extractions should be carried out from the most ancient to most recent samples, to prevent modern DNA inadvertently being carried to ancient samples. The amount of sediment used in DNA extractions should capture a representative picture of the biota present in a sample. Despite suggestions that bulk DNA extractions from up to 10 g of material can improve detection of taxa and better represent the diversity of the area of interest (e.g., Taberlet et al., 2012b; Coolen et al., 2013), using such large volumes of sediment is often not practical and can be quite costly in this field where typically many samples are processed. Instead, numerous studies have used replicate extractions of a smaller sample size (e.g., 0.25 g; Table 2) to increase the likelihood of yielding aDNA from rare taxa, as well as successive DNA extractions from a single 0.25 g sediment sample (e.g., Willerslev et al., 2003). Post-extraction, the use of RNA-probe based enrichment approaches coupled with shotgun sequencing, a common technique in aDNA research, may furthermore drastically improve the detection of rare taxa (Horn, 2012).

While it would be ideal to find one extraction method that will yield the best quality data and enable standardisation across ancient marine sediment studies, the type of sediment or target organisms may require some adjustments of standard protocols (Hermans et al., 2018). Extraction methods can bias the diversity observed due to differential resilience of taxa to the cell-lysis method (Zhou et al., 1996; Young et al., 2015) and DNA binding capacities of different soil and sediment types (Lorenz and Wackernagel, 1994; Miller et al., 1999). As a result, the aDNA extraction efficiency can be poor and the detection of an aDNA signal lost. To date, a variety of commercial kits have been successfully used to isolate aDNA from sediments (Table 2). To further increase the yield of very low amounts of highly fragmented aDNA several studies have been utilising extraction protocols that include a liquid-silica DNA binding step (e.g., Brotherton et al., 2013 and Weyrich et al., 2017 for non-sediment samples) or ethylenediaminetetraacetic acid (EDTA) cell-lysis step (Slon et al., 2017; utilising cave-sediment samples). Other studies have replaced the Bead Solution in the DNeasy extraction kits (Qiagen; Table 2) by 1 M sodium phosphate

pH 9–10 and 15 vol% ethanol to efficiently release clay-adsorbed DNA, and to prevent DNA released from intact cells from adsorbing to clay minerals during the extraction (Direito et al., 2012; Orsi et al., 2017; More et al., 2018). The latter is especially important when working with low organic, high carbonate rocks and sediments (Direito et al., 2012).

Two points are particularly important to prevent contamination during extractions. Firstly, as with the samples themselves, it is crucial that all tools and reagents undergo rigorous sterilisation procedures before utilisation, such as by bleach and UV treatment of any packing material before entering ancient DNA facilities. Secondly, blank controls should be included for every step of the laboratory process, i.e., extraction/library preparation blank controls, sequencing and bioinformatic analysis controls (Ficetola et al., 2016). Controlling and monitoring contamination is particularly important when analysing bacterial diversity due to their presence in all laboratory environments and reagents (Weyrich et al., 2015). Optimally, extraction blanks are included in a 1:5 ratio (Willerslev and Cooper, 2005), with a bare minimum of one control with each set of extractions. Aside from bioinformatically removing any organisms determined in such extraction blanks from the investigated sample material, the contaminants should be tracked within a laboratory, and contaminant lists published alongside the data for reasons of data transparency and authenticity.

Post-extraction, many marine *sedaDNA* studies have employed methods that are routinely used for modern marine DNA analysis. Although modern DNA work is not exempt from precautions, there are several issues with aDNA work: (i) as outlined in Sections 2.1 and 2.2. aDNA is highly fragmented and degraded and any small amount of modern DNA present in the sample (from reagents, labs or living cells that were present in the sediment sample) will amplify over the aDNA; (ii) sampling and extraction controls are often not included in the sequencing sample; (iii) PCRs are often inhibited due to the co-extraction of humic substances, pigments and heavy metals along with DNA (Webster et al., 2003 and references therein), requiring adequate removal of these impurities (e.g., Coolen et al., 2009); (iv) successful PCRs are prone to bias due to random amplification in reactions that contain very low amounts of DNA template, thus PCR drift (stochastic variation in the first PCR cycles) can occur (Wagner et al., 1994; Polz and Cavanaugh, 1998; Webster et al., 2003). More importantly, the number of, e.g., bacterial 16S and eukaryotic 18S rRNA operons can greatly vary between per genome and per cell and can cause a biased representation of the past community structure (e.g., Klappenbach et al., 2001). The above biases can be reduced and the detection limit lowered when PCR approaches selectively, amplifying particular groups of organisms indicative of environmental changes, are paired with independent geochemical proxies (e.g., Coolen et al., 2004, 2006, 2009). However, we strongly advocate for the use of strict aDNA methodologies and facilities in order to achieve the generation of authentic marine *sedaDNA* data, following the guidelines in this review.

Shotgun metagenomics are currently widely accepted and the least biased method to analyse the broad diversity of ancient environmental samples (e.g., Slon et al., 2017). Although only a small portion of the generated sequence data might be attributable to the ancient organism in question (Morard et al., 2017), next generation sequencing (NGS) generates large quantities of data that enable meaningful statistics, with the additional benefit of preserving the relative proportion of detected taxa. To analyse aDNA sequence data, robust bioinformatic pipelines (e.g., Paleomix, Schubert et al., 2014) have been developed and are available for the application to marine *sedaDNA*, integrating damage detection algorithms (e.g., Ginolhac et al., 2011; Kistler et al., 2017) that enable the distinction between ancient and modern signals. Determining the extent of cytosine residues deamination (C to T and G to A, Weyrich et al., 2017) should also be considered to assess authenticity of aDNA sequences, especially when the data was generated from mixed communities, such as from marine *sedaDNA*. It is furthermore crucial to carefully screen sequencing data for any low-complexity reads, which

may get incorrectly assigned to taxa during alignments against genetic databases, as well as ensuring that taxonomic assignments in the database of choice are correct. Bioinformatic pipelines removing such misidentification-derived errors do not currently exist and should be the focus of future research, as well as the comparison of shotgun and amplicon marine *sedaDNA* data to accurately determine biases and analysis strategies best suited to this new discipline.

4. Future marine *sedaDNA* sampling considerations

4.1. Equipment and installations required aboard IODP platforms

In addition to the recent upgrades and investments IODP has made to enable sediment sampling suitable for Biosphere Frontiers theme (Section 3.1) we suggest the following items to facilitate contamination-free sediment sampling and the tracing of contaminants.

(i) Laboratories in which sampling for *sedaDNA* is undertaken should be carefully chosen to minimise contamination. Rapid transport of the core from the deck to the lab, thorough decontamination measures (see Section 3.4), and easy access to fridges or freezers are crucial. While a positively air-pressured lab (standard for aDNA laboratories) may not be feasible, air-flow can be reduced by keeping all doors shut and fans off during aDNA sampling. Contamination by human DNA from analysts can be greatly reduced by wearing adequate protective clothing (gloves, facemask, freshly laundered/disposable lab coat/overall). A detailed record or any molecular work undertaken in ship-board labs should be maintained by IODP, and under no circumstances should *sedaDNA* sampling be conducted in labs used previously to run PCRs (see Section 3.4). Alternatively, the equipment of a shipping container exclusively dedicated to *sedaDNA* sampling could be a good solution to spatially separate *sedaDNA* sampling aboard drilling-platforms and installation could be as required during expeditions that involve *sedaDNA* sampling.

(ii) DNA is likely to behave quite different from chemical tracers and microspheres currently used to track contamination. With constantly advancing technologies in the field of synthetic biology, the possibility arises to develop ‘non-biological DNA’ with known sequences. Such non-viable DNA tags are already used in the oil industry, where a different tag is introduced into oil pipes monthly to monitor when and where leaks occur (Forecast Technology Ltd). Using such tags during seafloor coring operations instead of chemical tracers should enable a precise assessment of contamination by environmental DNA, where bioinformatics pipelines could be adjusted to detect and quantify the amount of tags present in the final sequencing data.

4.2. Ground-truthing marine *sedaDNA* research and data

To ground-truth marine *sedaDNA* studies and to ensure the generation of authentic *sedaDNA* data we suggest future research in this field to focus on the following aspects:

- (i) The establishment of a public record of common contaminants. This can be achieved, for example, through an inter-lab comparison focused on analysing the same samples and integrating extraction blanks to trace contaminants associated with particular coring equipment, ship- and land- based laboratories.
- (ii) Investigation of factors that might considerably bias marine *sedaDNA* data. This might include information on sediment-type and environmental condition dependent *sedaDNA* preservation, taxon-specific DNA degradation rates, average *sedaDNA* fragment length, and shotgun and amplification-based aDNA data comparisons.
- (iii) Ongoing enrichment of genetic reference databases for modern marine plankton, to enable taxonomic assignment of the hundreds of thousands of ancient sequences expected to be found in marine sediments.

- (iv) The inclusion of negative controls during extractions, library preparations and in sequencing runs, and the publication of the results in the context of independent multiproxy biological and environmental metadata obtained from the same sediment interval.
- (v) Once (i) - (iv) are addressed, the development of a dedicated *seDaDNA* coring proposal is encouraged, in which sediment cores are collected using the above outlined, best-suited coring strategies, sampling and analysis procedures. During such an expedition, basic questions such as optimal on-board contamination tracing techniques, feasible work-flows, spatial replication required to achieve representative community data, and age to which marine *seDaDNA* can be determined should be addressed. Such baseline data is missing to date and remains the most important step towards the generation of authentic aDNA data from marine sediments.

5. Application of marine *seDaDNA* research guidelines to other contamination susceptible environments

5.1. Permafrost

Permafrost molecular biological studies provide the opportunity to study living organisms that have successfully adapted to extremely cold environments and comprise an analogous cryogenic environment to that found on other planets, such as Mars (Amato et al., 2010). Molecular investigations have focussed on humans (Rasmussen et al., 2010), plants (Willerslev et al., 2003), megafauna (Boessenkool et al., 2012), fungi (Bellemain et al., 2013) and microbes (Willerslev et al., 2004). Permafrost top layers are characterised by a more abundant and diverse microbial community compared to the deeper soil (Gittel et al., 2014). To overcome the hurdle of distinguishing between the modern and ancient DNA signal, metatranscriptomics have been applied to identify the active community only (e.g., Coolen and Orsi, 2015). Despite the challenges in experimental approaches, such as rapid community shifts after thawing even at nearly ambient conditions (Negandhi et al., 2016), studies of permafrost environments have advanced our understanding of feedback loops associated with the response of extremophiles to warming, ultimately informing modelling studies including marine palaeo-environments.

Sampling for ice and permafrost in polar regions is challenging in terms of logistics and minimising contamination risks for both the sample and the sampled environment. For example, permafrost soil samples are, like marine sediment cores, retrieved through drilling, which can introduce microbial contaminants to the deeper permafrost soil layers as the drill head and liquid pass through the top active soil layer (Bang-Andreasen et al., 2017). Additionally, the cryosphere has been accumulating industrial chemicals and metals since the 1850's (McConnell et al., 2007), so that the present-day microbial community is now capable of degrading industrial contaminants, thereby representing an anthropogenically-adapted rather than an original pristine community (Hauptmann et al., 2017). With both these newly adapted anthropogenic and drilling fluid communities containing characteristics for heavy metal degradation, distinguishing indigenous ice core or permafrost communities from drilling fluid communities will become more difficult in the future (Miteva et al., 2014). Therefore, the described guidelines in this review for distinguishing ancient from modern and contaminant signals, as well as the need for aseptic sampling procedures, are highly applicable to permafrost environments and, more generally, the cryosphere.

5.2. Planetary exploration

The methodologies advocated in this review that enable aDNA in marine sediments to be distinguished from modern DNA are also applicable to the search for life on other planets or moons. Astrobiologists are especially interested in the possibility of detection of Life 2.0, where

the life has an independent genesis to that on Earth. The search for life beyond Earth has been potentially possible since the 1970s, with the two Viking lander missions to Mars, but there are other possible targets in our solar system, notably some of the moons around Jupiter and Saturn (e.g., Europa, Titan). Space technology has now reached the point where the detection of life, if it exists or existed elsewhere in the solar system, is becoming a realistic possibility in the next 50 years. There have been several rovers that have carried out successful exploration of the surface of Mars, including Curiosity, the Mars Science Laboratory that in 2018 is mid-way through its predicted mission (Grotzinger et al., 2014). The rover Mars 2020 is being designed at present to test for evidence of life in the near-surface environment. It will drill, collect and cache samples from the Martian surface, which will then be returned to Earth for more detailed analysis (Beatty et al., 2015). Sample return from Mars to Earth is planned for the end of the 2020's (Foust, 2018). Active planning is also ongoing for possible missions to land and analyse materials from the surfaces of moons such as Europa and Titan, by both NASA and the European Space Agency. For example, Europa (a moon of Jupiter) is known to have a global saltwater ocean below its icy crust, as well as a rocky seafloor, so is one of the highest priority targets in the search for present-day life beyond Earth (Hand et al., 2017). A key concern with this solar system exploration is planetary protection, which is governed by the United Nations Outer Space Treaty (United Nations Office for Disarmament Affairs, 2015) and the Committee on Space Research (COSPAR) of the International Committee for Science. There are two important categories of planetary protection. The first is "forward contamination", where Earth-derived microbial life hitches a ride on spacecraft and contaminates parts of a planetary surface being explored. The second is "backward contamination", where life from an explored planet or moon is inadvertently returned to Earth, maybe in a spacecraft or within a rock sample. The relevance to aDNA analytical protocols is in forward contamination (i.e., the risk of contaminating sample material that could lead to data misinterpretations, and/or generally introducing Earth contaminants to other planets; Rummel and Conley, 2017). It should be noted that if indeed there is or was life on other planetary bodies, it may well not be based on a genetic code composed of DNA and RNA. Independently originated Life 2.0 would be highly unlikely to have evolved exactly the same nucleic acid genetic code as life on Earth (e.g., Rummel and Conley, 2017). Indeed, it has been postulated that an alternative biosphere could exist as a "shadow biosphere" on Earth (Davies et al., 2009). If DNA or RNA-based extant life is found on Mars, for example, then it is most likely that it would represent either past natural exchange of rocks between the two planets (panspermia), or anthropogenic forward contamination. Therefore, the procedures used for distinguishing indigenous life in planetary exploration will need broadening to include the possibility of life with a different genetic code. The protocols developed for aDNA sampling of marine sediments on Earth, including the ability to distinguish from modern DNA, have relevance for the designing of methods to look for past life on Mars or outer solar system moons using molecular biology techniques (Beatty et al., 2015; Hand et al., 2017).

6. Conclusions

Ancient DNA in marine deep-sea sediments holds the potential to open a new era of marine palaeo-environment and -climate reconstruction. However, anti-contamination measures central to all aDNA research have logistical constraints and are particularly poorly-suited to shipboard sediment sampling and processing. For example, sterile coring equipment and ultra-clean laboratories are usually not available on any type of drilling platform. Current and future IODP drilling vessels are aware of the increasing need for improved and innovative solutions to coring, non-contaminant drill fluids and appropriate laboratories and storage facilities. Such logistical advances should go hand-in-hand with the establishment of new criteria and

standards to ensure the acquisition and preservation of sediment cores with minimal environmental contaminants. Complementary genetic and geochemical information currently available to date suggests that, realistically, environmental reconstructions based on marine *sedDNA* from past plankton can be achieved for at least the last glacial-interglacial cycle, and potentially back to ~400,000 years. These guidelines can be applied in other scientific areas to facilitate and optimise research conducted in extremely remote locations, contamination-susceptible environmental samples, and even during the future exploration of other planets.

Acknowledgements

This work is the outcome of a workshop on “Ancient DNA in marine sediments”, held at Macquarie University, Australia, in December 2017. We thank Prof. Ian Paulsen for insightful discussions around the topic, and Mrs. Kelly-Anne Lawler and Ms. Louise Fleck for their support with the organisation and running of the workshop.

Funding sources

The above-named workshop was supported by the Macquarie University Marine Research Centre (MQMarine), the Department of Biological Sciences, and the Faculty of Science and Engineering, Macquarie University, North Ryde, Australia. LA was supported by the Australia and New Zealand IODP Consortium (ANZIC) 2017 Special Analytical Funding and MQMarine. MJLC was supported by Australian Research Council's (ARC) Discovery Projects DP160102587 and DP18100982 as well as by ANZIC (IODP Exp. 364 post cruise funding). LKA and SCG were supported by the ARC LIEF funding scheme LE160100067 provided to ANZIC. MO was supported by an ARC Discovery Project DP15012326. MO and AF were supported by an ARC Laureate Fellowship FL14010021 to Ian Paulsen. MS acknowledges the support of ARC Linkage Project LP160100839. JWM acknowledges the support of ARC Discovery Project DP110103668. AC is supported by ARC Laureate Fellowship (FL140100260) and LSW by ARC Future Fellowship (FT180100407), both AC and LSW are also supported by the ARC Centre of Excellence for Australian Biodiversity and Heritage (CABAH; 170100015CE). NRF was funded by the University of Adelaide RTP scholarship. FL is supported by an Early Postdoc Mobility fellowship from the Swiss National Science Foundation.

References

- Expedition 330 Scientists, 2012. Methods. In: Koppers, A.A.P., Yamazaki, T., Geldmacher, J. (Eds.), *The Expedition 330 Scientists*, Proc. IODP. 330 Integrated Ocean Drilling Program Management International, Inc., Tokyo. <https://doi.org/10.2204/iodp.proc.330.102.2012>.
- Sun, Z., Jian, Z., Stock, J.M., Larsen, H.C., Klaus, A., Alvarez Zarikian, C.A., Expedition 367/368 Scientists, 2018. *South China Sea Rifted Margin*. Proceedings of the International Ocean Discovery Program, 367/368. International Ocean Discovery Program, College Station, TX. <https://doi.org/10.14379/iodp.proc.367368.2018>.
- United Nations Office for Disarmament Affairs, 2015. Treaty on Principles Governing the Activities of States in the Exploration and Use of Outer Space, including the Moon and Other Celestial Bodies, Article IX, version 2015, signed by almost all nation states, including all the current and aspiring space-faring nation states. Available online at: http://disarmament.un.org/treaties/o/outer_space.
- Van Everdingen, R.O. (Ed.), 1998. Multi-Language Glossary of Permafrost and Related Ground-Ice Terms in Chinese, English, French, German, Icelandic, Italian, Norwegian, Polish, Romanian, Russian, Spanish, and Swedish. International Permafrost Association, Terminology Working Group.
- Alawi, M., Schneider, B., Kallmeyer, J., 2014. A procedure for separate recovery of extra- and intracellular DNA from a single marine sediment sample. *J. Microbiol. Methods* 104, 36–42.
- Allentoft, M.E., Collins, M., Harker, D., Haile, J., Oskam, C.L., Hale, M.L., Campos, P.F., Samaniego, J.A., Gilbert, M.T.P., Willerslev, E., Zhang, G., 2012. The half-life of DNA in bone: measuring decay kinetics in 158 dated fossils. In: Proceedings of the Royal Society of London B: Biological Sciences, (p.rspb20121745).
- Alsos, I.G., Sjögren, P., Edwards, M.E., Landvik, J.Y., Gielly, L., Forwick, M., Coissac, E., Brown, A.G., Jakobsen, L.V., Foreid, M.K., Pedersen, M.W., 2015. Sedimentary ancient DNA from Lake Skartjørna, Svalbard: Assessing the resilience of arctic flora to Holocene climate change. *The Holocene* 26, 627–642.
- Amato, P., Doyle, S.M., Battista, J.R., Christner, B.C., 2010. Implications of subzero metabolic activity on long-term microbial survival in terrestrial and extraterrestrial permafrost. *Astrobiology* 10, 789–798.
- Armbricht, L.H., Lowe, V., Escutia, C., Iwai, M., McKay, R., Armand, L.K., 2018. Variability in diatom and silicoflagellate assemblages during mid-Pliocene glacial-interglacial cycles determined in Hole U1361A of IODP Expedition 318, Antarctic Wilkes Land margin. *Mar. Micropaleontol.* 139, 28–41.
- Armbricht, L., Herrando-Perez, S., Eisenhofer, R., Hallegraeff, G., Bolch, C., Cooper, A. An optimised method for the extraction of ancient eukaryotic DNA from marine sediments. In prep.
- Bang-Andreasen, T., Schostag, M., Priemé, A., Elberling, B., Jacobsen, C.S., 2017. Potential microbial contamination during sampling of permafrost soil assessed by tracers. *Sci. Rep.* 7, 43338.
- Beatty, D.W., Hays, L.E., Williford, K., Farley, K., 2015. Sample Science Input to Landing Site selection for Mars 2020: an In-Situ Exploration and Sample Caching Rover. *Meteorit. Planet. Sci.* 50 (S1) (Art-No. 5340).
- Bellemain, E., Davey, M.L., Kauserud, H., Epp, L.S., Boessenkool, S., Coissac, E., Geml, J., Edwards, M., Willerslev, E., Gussarova, G., Taberlet, P., 2013. Fungal palaeodiversity revealed using high-throughput metabarcoding of ancient DNA from arctic permafrost. *Environ. Microbiol.* 15, 1176–1189.
- Bernhard, J.M., Bowser, S.S., 1999. Benthic foraminifera of dysoxic sediments: chloroplast sequestration and functional morphology. *Earth Sci. Rev.* 46, 149–165.
- Bickle, M., Arculus, R., Barrett, P., DeConto, R., Camoin, G., Edwards, K., Fisher, F., et al., 2011. Illuminating Earth's Past, Present and Future—The Science Plan for the International Ocean Discovery Program 2013–2023. Integrated Ocean Drilling Program, Washington, DC Available at: <http://www.iodp.org/about-iodp/iodp-science-plan-2013-2023>.
- Bidle, K.D., Lee, S., Marchant, D.R., Falkowski, P.G., 2007. Fossil genes and microbes in the oldest ice on Earth. *Proc. Natl. Acad. Sci. U. S. A.* 104, 13455–13460.
- Bienhold, C., Zinger, L., Boetius, A., Ramette, A., 2016. Diversity and biogeography of bathyal and abyssal seafloor bacteria. *PLoS One* 11, e0148016.
- Bissett, A., Gibson, J.A.E., Jarman, S.N., Swadling, K.M., Cromer, L., 2005. Isolation, amplification, and identification of ancient copepod DNA from lake sediments. *Limnol. Oceanogr. Methods* 3, 533–542.
- Blaser, M.J., Cardon, Z.G., Cho, M.K., Dangl, J.L., Donohue, T.J., Green, J.L., Knight, R., Maxon, M.E., Northen, T.R., Pollard, K.S., Brodie, E.L., 2016. Toward a predictive understanding of Earth's microbiomes to address 21st century challenges. *mBio* 7 (3), e00714–e00716.
- Boere, A.C., Abbas, B., Rijpstra, W.I.C., Versteegh, G.J.M., Volkman, J.K., Damsté, J.S.S., Coolen, M.J.L., 2009. Late-Holocene succession of dinoflagellates in an Antarctic fjord using a multi-proxy approach: paleoenvironmental genomics, lipid biomarkers and palynomorphs. *Geobiology* 7, 265–281.
- Boere, A.C., Rijpstra, W.I.C., De Lange, G.J., Sinninghe Damsté, J.S., Coolen, M.J.L., 2011. Preservation potential of ancient plankton DNA in Pleistocene marine sediments. *Geobiology* 9, 377–393.
- Boessenkool, S., Epp, L.S., Haile, J., Bellemain, E.V.A., Edwards, M., Coissac, E., Willerslev, E., Brochmann, C., 2012. Blocking human contaminant DNA during PCR allows amplification of rare mammal species from sedimentary ancient DNA. *Mol. Ecol.* 21, 1806–1815.
- Boudreau, B.P., 1998. Mean mixed depth of sediments: the wherefore and the why. *Limnol. Oceanogr.* 43, 524–526.
- Briggs, A.W., Stenzel, U., Johnson, P.L., Green, R.E., Kelso, J., Prüfer, K., Meyer, M., Krause, J., Ronan, M.T., Lachmann, M., Pääbo, S., 2007. Patterns of damage in genomic DNA sequences from a Neandertal. *Proc. Natl. Acad. Sci.* 104, 14616–14621.
- Brooks, J.J., Grice, K., 2011. Biomarkers (Molecular Fossils). In: Reitner, J., Thiel, V., Dordrecht (Eds.), *Encyclopedia of Geobiology*. Springer, Netherlands, pp. 147–167.
- Brotherton, P., Haak, W., Templeton, J., Brandt, G., Soubrier, J., Adler, C.J., Richards, S.M., Der Sarkissian, C., Ganslmeier, R., Friederich, S., Dresely, V., 2013. Neolithic mitochondrial haplogroup H genomes and the genetic origins of Europeans. *Nat. Commun.* 4, 1764.
- Calvert, S.E., Vogel, J.S., Southon, J.R., 1987. Carbon accumulation rates and the origin of the Holocene sapropel in the Black Sea. *Geology* 15, 918–921.
- Carradec, Q., Pelletier, E., Da Silva, C., Alberti, A., Seeluthner, Y., Blanc-Mathieu, R., Lima-Mendez, G., Rocha, F., Tirichine, L., Labadie, K., Kirilovsky, A., 2018. A global atlas of eukaryotic genes. *Nat. Commun.* 9, 373.
- Ceccherini, M.T., Ascher, J., Pietramellara, G., Vogel, T.M., Nannipieri, P., 2007. Vertical advection of extracellular DNA by water capillarity in soil columns. *Soil Biol. Biochem.* 39, 158–163.
- Cockell, C.S., Coolen, M., Schaefer, B., Grice, K., Gulick, S.P.S., Morgan, J.V., Kring, D.A., Osinski, G., 2017. Deep Subsurface Microbial Communities shaped by the Chicxulub Impactor. In: American Geophysical Union, Fall Meeting 2017, Abstract #P23H-06, <http://adsabs.harvard.edu/abs/2017AGUFM.P23H.06C>.
- Coolen, M.J., 2011. 7000 years of *Emiliania huxleyi* viruses in the Black Sea. *Science* 333, 451–452.
- Coolen, M.J., Orsi, W.D., 2015. The transcriptional response of microbial communities in thawing Alaskan permafrost soils. *Front. Microbiol.* 6, 197.
- Coolen, M.J., Overmann, J., 1998. Analysis of subsolfid molecular remains of purple sulfur bacteria in a lake sediment. *Appl. Environ. Microbiol.* 64, 4513–4521.
- Coolen, M.J., Overmann, J., 2007. 217 000-year-old DNA sequences of green sulfur bacteria in Mediterranean sapropels and their implications for the reconstruction of the paleoenvironment. *Environ. Microbiol.* 9, 238–249.
- Coolen, M.J., Muiyzer, G., Rijpstra, W.I.C., Schouten, S., Volkman, J.K., Damsté, J.S.S., 2004. Combined DNA and lipid analyses of sediments reveal changes in Holocene haptophyte and diatom populations in an Antarctic lake. *Earth Planet. Sci. Lett.* 223, 225–239.
- Coolen, M.J., Boere, A., Abbas, B., Baas, M., Wakeham, S.G., Sinninghe Damsté, J.S., 2006. Ancient DNA derived from alkenone-biosynthesizing haptophytes and other algae in Holocene sediments from the Black Sea. *Paleoceanography* 21, PA1005.
- Coolen, M.J.L., Talbot, H.M., Abbas, B.A., Ward, C., Schouten, S., Volkman, J.K., Damsté, J.S.S., 2008. Sources for sedimentary bacteriophanepolys as revealed by 16S rDNA stratigraphy. *Environ. Microbiol.* 10, 1783–1803.
- Coolen, M.J.L., Saenz, J.P., Giosan, L., Trowbridge, N.Y., Dimitrov, P., Dimitrov, D.,

- Eglinton, T.I., 2009. DNA and lipid molecular stratigraphic records of haptophyte succession in the Black Sea during the Holocene. *Earth Planet. Sci. Lett.* 284, 610–621.
- Coolen, M.J., Orsi, W.D., Balkema, C., Quince, C., Harris, K., Sylva, S.P., Filipova-Marinova, M., Giosan, L., 2013. Evolution of the plankton paleome in the Black Sea from the Deglacial to Anthropocene. *Proc. Natl. Acad. Sci.* 110, 8609–8614.
- Cooper, A., Poinar, H.N., 2000. Ancient DNA: do it right or not at all. *Science* 289, 1139.
- Corinaldesi, C., Danovaro, R., Dell'Anno, A., 2005. Simultaneous recovery of extracellular and intracellular DNA suitable for molecular studies from marine sediments. *Appl. Environ. Microbiol.* 71, 46–50.
- Corinaldesi, C., Beolchini, F., Dell'Anno, A., 2008. Damage and degradation rates of extracellular DNA in marine sediments: implications for the preservation of gene sequences. *Mol. Ecol.* 17, 3939–3951.
- Corinaldesi, C., Barucca, M., Luna, G.M., Dell'Anno, A., 2011. Preservation, origin and genetic imprint of extracellular DNA in permanently anoxic deep-sea sediments. *Mol. Ecol.* 20, 642–654.
- Corinaldesi, C., Tangherlini, M., Luna, G.M., Dell'Anno, A., 2014. Extracellular DNA can preserve the genetic signatures of present and past viral infection events in deep hypersaline anoxic basins. *Proc. R. Soc. Lond. B Biol. Sci.* 281, 20133299.
- Corinaldesi, C., Tangherlini, M., Manea, E., Dell'Anno, A., 2018. Extracellular DNA as a genetic recorder of microbial diversity in benthic deep-sea ecosystems. *Sci. Rep.* 8, 1839.
- Cox, T.L., Gan, H.M., Moreau, J.W., 2019. Seawater recirculation through subducting sediments sustains a deeply buried population of sulfate-reducing bacteria. *Geobiology* 17, 172–184.
- D'Andrea, W.J., Lage, M., Martiny, J.B.H., Laatsch, A.D., Amaral-Zettler, L.A., Sogin, M.L., Huang, Y., 2006. Alkenone producers inferred from well-preserved 18S rDNA in Greenland lake sediments. *J. Geophys. Res.* 111, G0313.
- Davies, P.C., Benner, S.A., Cleland, C.E., Lineweaver, C.H., McKay, C.P., Wolfe-Simon, F., 2009. Signatures of a shadow biosphere. *Astrobiology* 9, 241–249.
- De Vargas, C., Audic, S., Henry, N., Decelle, J., Mahé, F., Logares, R., Lara, E., Berney, C., Le Bescot, N., Probert, I., Carmichael, M., 2015. Eukaryotic plankton diversity in the sunlit ocean. *Science* 348, 1261605.
- Dell'Anno, A., Stefano, B., Danovaro, R., 2002. Quantification, base composition, and fate of extracellular DNA in marine sediments. *Limnol. Oceanogr.* 47, 899–905.
- Dell'Anno, A., Corinaldesi, C., Anno, A.D., 2004. Degradation and turnover of extracellular DNA in marine sediments: ecological and methodological considerations. *Appl. Environ. Microbiol.* 70, 4384–4386.
- Demanèche, S., Bertolla, F., Buret, F., Nalin, R., Sailland, A., Auriol, P., Vogel, T.M., Simonet, P., 2001. Laboratory-scale evidence for lightning-mediated gene transfer in soil. *Appl. Environ. Microbiol.* 67, 3440–3444.
- Direito, S.O., Marees, A., Röling, W.F., 2012. Sensitive life detection strategies for low-biomass environments: optimizing extraction of nucleic acids adsorbing to terrestrial and Mars analog minerals. *FEMS Microbiol. Ecol.* 81, 111–123.
- Feek, D.T., Horrocks, M., Baisden, W.T., Flenley, J., 2011. The Mk II sampler: a device to collect sediment cores for analysis of uncontaminated DNA. *J. Paleolimnol.* 45, 115–119.
- Ficetola, G.F., Taberlet, P., Coissac, E., 2016. How to limit false positives in environmental DNA and metabarcoding? *Mol. Ecol. Resour.* 16, 604–607.
- Foust, J., 2018. ESA awards Mars sample return study contracts as international cooperation plans take shape. *SpaceNews* (July 8), 2018. Available at: <https://spacenews.com/esa-awards-mars-sample-return-study-contracts-as-international-cooperation-plans-take-shape/>.
- Früh-Green, G., Orcutt, B., Green, S., 2015. Expedition 357 scientific prospectus: Atlantis Massif serpentinization and life. *Int. Ocean Discov. Prog.* <https://doi.org/10.14379/iodp.proc.327.2015>.
- Fry, J.C., Parkes, R.J., Cragg, B.A., Weightman, A.J., Webster, G., 2008. Prokaryotic biodiversity and activity in the deep subsurface biosphere. *FEMS Microbiol. Ecol.* 66, 181–196.
- Fryer, P., Wheat, C.G., Williams, T., Albers, E., Bekins, B., Debret, B.P.R., Deng, J., Dong, Y., Eickenbusch, P., Frery, E.A., Ichiyama, Y., Johnson, K., Johnston, R.M., Kevorkian, R.T., Kurz, W., Magalhaes, V., Mantovanelli, S.S., Menapace, W., Menzies, C.D., Michibayashi, K., Moyer, C.L., Mullane, K.K., Park, J.-W., Price, R.E., Ryan, J.G., Shervais, J.W., Sissmann, O.J., Suzuki, S., Takai, K., Walter, B., Zhang, R., 2018. Expedition 366 methods. In: *Proceedings of the International Ocean Discovery Program*. vol. 366 <https://doi.org/10.14379/iodp.proc.366.102.2018.publications.iodp.org>.
- Furukawa, Y., Horiuchi, M., Kakegawa, T., 2013. Selective stabilization of ribose by borate. *Orig. Life Evol. Biosph.* 43, 353–361.
- Giguet-Covex, C., Pansu, J., Arnaud, F., Rey, P.J., Griggo, C., Gielly, L., Domaizon, I., Coissac, E., David, F., Choler, P., Poulenard, J., Taberlet, P., 2014. Long livestock farming history and human landscape shaping revealed by lake sediment DNA. *Nat. Commun.* 5, 3211.
- Ginolhac, A., Rasmussen, M., Gilbert, M.T.P., Willerslev, E., Orlando, L., 2011. mapDamage: testing for damage patterns in ancient DNA sequences. *Bioinformatics* 27, 2153–2155.
- Giosan, L., Coolen, M.J., Kaplan, J.O., Constantinescu, S., Filip, F., Filipova-Marinova, M., Kettner, A.J., Thom, N., 2012. Early anthropogenic transformation of the Danube-Black Sea system. *Sci. Rep.* 2 (582).
- Gittel, A., Bárta, J., Kohoutova, I., Schnecker, J., Wild, B., Čapek, P., Kaiser, C., Torsvik, V.L., Richter, A., Schleper, C., Ulrich, T., 2014. Site- and horizon-specific patterns of microbial community structure and enzyme activities in permafrost-affected soils of Greenland. *Front. Microbiol.* 5, 541.
- Grotzinger, J.P., Sumner, D.Y., Kah, L.C., Stack, K., Gupta, S., Edgar, L., Rubin, D., Lewis, K., Schieber, J., Mangold, N., Milliken, R., Conrad, P.G., DesMarais, D., Farmer, J., Siebach, K., Calef, F., Hurowitz, J., McLennan, S.M., Ming, D., Vaniman, D., Crisp, J., Vasavada, A., Edgett, K.S., Malin, M., Blake, D., Gellert, R., Mahaffy, P., Wiens, R.C., Maurice, S., Grant, J.A., Wilson, S., Anderson, R.C., Beegle, L., Arvidson, R., Hallet, B., Sletten, R.S., Rice, M., Bell, J., Griffes, J., Ehlmann, B., Anderson, R.B., Bristow, T.F., Dietrich, W.E., Dromart, G., Eigenbrode, J., Fraeman, A., Hardgrove, C., Herkenhoff, K., Jandura, L., Kocurek, G., Lee, S., Leshin, L.A., Leveille, R., Limonadi, D., Maki, J., McCloskey, S., Meyer, M., Minitti, M., Newsom, H., Oehler, D., Okon, A., Palucis, M., Parker, T., Rowland, S., Schmidt, M., Squyres, S., Steele, A., Stolper, E., Summons, R., Treiman, A., Williams, R., Yingst, A., 2014. A habitable fluvio-lacustrine environment at Yellowknife Bay, Gale Crater, Mars. *Science* 343, 1242777.
- Haile, J., 2012. Ancient DNA extraction from soils and sediments. In: Shapiro, B., Hofreiter, M. (Eds.), *Ancient DNA, Methods and Protocols* (2012). Springer New York, Dordrecht, Heidelberg, London, pp. 57–64.
- Haile, J., Holdaway, R., Oliver, K., Bunce, M., Gilbert, M.T., Nielsen, R., Munch, K., Ho, S.Y., Shapiro, B., Willerslev, E., 2007. Ancient DNA chronology within sediment deposits: are paleobiological reconstructions possible and is DNA leaching a factor? *Mol. Biol. Evol.* 24, 982–989.
- Hand, K.P., Murray, A.E., Garvin, J.B., Brinkerhoff, W.B., Christner, B.C., Edgett, K.S., Ehlmann, B.L., German, C.R., Hayes, A.G., Hoehler, T.M., Horst, S.M., Lumine, J.J., Nealson, K.H., Parancas, C., Schmidt, B.E., Smith, D.E., Rhoden, A.R., Russell, M.J., Tomperton, A.S., Willis, P.A., Yingst, R.A., Phillips, C.B., Cable, M.L., Craft, K.L., Hofmann, A.E., Nordheim, T.A., Pappalardo, R.P., Project Engineering Team, 2017. Report of the Europa Lander Science Definition Team. National Aeronautics and Space Administration.
- Hauptmann, A.L., Sichert-Pontén, T., Cameron, K.A., Bælum, J., Plichta, D.R., Dalgaard, M., Stibal, M., 2017. Contamination of the Arctic reflected in microbial metagenomes from the Greenland ice sheet. *Environ. Res. Lett.* 12, 074019.
- Hay, B.J., 1988. Sediment accumulation in the central western Black Sea over the past 5,100 years. *Paleoceanography* 3, 491–508.
- Hay, B.J., Arthur, M.A., Dean, W.E., Neff, E.D., Honjo, S., 1991. Sediment deposition in the Late Holocene abyssal Black Sea with climatic and chronological implications. In: *Deep Sea Res. Part A. Oceanogr. Res. Pap.* 38, pp. S1211–S1235.
- Hermans, S.M., Buckley, H.L., Lear, G., 2018. Optimal extraction methods for the simultaneous analysis of DNA from diverse organisms and sample types. *Mol. Ecol. Resour.* 18, 557–569.
- Horn, S., 2012. Target enrichment via DNA hybridization capture. In: Shapiro, B., Hofreiter, M. (Eds.), *Ancient DNA, Methods and Protocols* (2012). Springer New York, Dordrecht, Heidelberg, London, pp. 177–188.
- Hou, W., Dong, H., Li, G., Yang, J., Coolen, M.J., Liu, X., Wang, S., Jiang, H., Wu, X., Xiao, H., Lian, B., 2014. Identification of photosynthetic plankton communities using sedimentary ancient DNA and their response to late-Holocene climate change on the Tibetan Plateau. *Sci. Rep.* 4, 6648.
- Inagaki, F., Hinrichs, K.U., Kubo, Y., Bowles, M.W., Heuer, V.B., Hong, W.L., Hoshino, T., Ijiri, A., Imachi, H., Ito, M., Kaneko, M., 2015. Exploring deep microbial life in coal-bearing sediment down to ~2.5 km below the ocean floor. *Science* 349, 420–424.
- Jónsson, H., Ginolhac, A., Schubert, M., Johnson, P.L., Orlando, L., 2013. mapDamage2.0: fast approximate Bayesian estimates of ancient DNA damage parameters. *Bioinformatics* 29, 1682–1684.
- Jørgensen, T., Kjaer, K.H., Haile, J., Rasmussen, M., Boessens, S., Andersen, K., Coissac, E., Taberlet, P., Brochmann, C., Orlando, L., Gilbert, M.T., Willerslev, E., 2012. Islands in the ice: detecting past vegetation on Greenlandic nunataks using historical records and sedimentary ancient DNA meta-barcoding. *Mol. Ecol.* 21, 1980–1988.
- Kirkpatrick, J.B., Walsh, E.A., D'Hondt, S., 2016. Fossil DNA persistence and decay in marine sediment over hundred-thousand-year to million-year time scales. *Geology* 44, 615–618.
- Kistler, L., Ware, R., Smith, O., Collins, M., Allaby, R.G., 2017. A new model for ancient DNA decay based on paleogenomic meta-analysis. *Nucleic Acids Res.* 45, 6310–6320.
- Klappenbach, J.A., Saxman, P.R., Cole, J.R., Schmidt, T.M., 2001. RnDb: the ribosomal RNA operon copy number database. *Nucleic Acids Res.* 29, 181–184.
- Klemetsen, T., Raknes, I.A., Fu, J., Agafonov, A., Balasundaram, S.V., Tartari, G., Robertsen, E., Willassen, N.P., 2017. The MAR databases: development and implementation of databases specific for marine metagenomics. *Nucleic Acids Res.* 46, D692–D699.
- Leite, D., Leitão, A., Schaan, A.P., Marinho, A.N., Souza, S., Rodrigues-Carvalho, C., Cardoso, F., Ribeiro-dos-Santos, Á., 2014. Paleogenetic studies in Guajajara skeletal remains, Maranhão state, Brazil. *J. Anthropol.* 729120.
- Lejzerowicz, F., Esling, P., Majewski, W., Szczuciński, W., Decelle, J., Obadia, C., Arbizu, P.M., Pawłowski, J., 2013. Ancient DNA complements microfossil record in deep-sea subsurface sediments. *Biol. Lett.* 9, 20130283.
- Lever, M.A., Alperin, M., Engelen, B., Inagaki, F., Nakagawa, S., Steinsbu, B.O., Teske, A., 2006. Trends in basalt and sediment core contamination during IODP Expedition 301. *Geomicrobiol. J.* 23, 517–530.
- Levy-Booth, D.J., Campbell, R.G., Gulden, R.H., Hart, M.M., Powell, J.R., Klironomos, J.N., Pails, P., Swanton, C.J., Trevors, J.T., Dunfield, K.E., 2007. Cycling of extracellular DNA in the soil environment. *Soil Biol. Biochem.* 39 (12), 2977–2991.
- Lindahl, T., 1993. Instability and decay of the primary structure of DNA. *Nature* 362, 709.
- Lorenz, M.G., Wackernagel, W., 1994. Bacterial gene transfer by natural genetic transformation in the environment. *Microbiol. Rev.* 58, 563–602.
- Loucaides, S., Van Cappellen, P., Roubex, V., Moriceau, B., Ragueneau, O., 2011. Controls on the recycling and preservation of biogenic silica from biomineralization to burial. *Silicon* 4, 7–22.
- Lyon, D.Y., Monier, J.M., Dupraz, S., Freissinet, C., Simonet, P., Vogel, T.M., 2010. Integrity and biological activity of DNA after UV exposure. *Astrobiology* 10, 285–292.
- Lyra, C., Sinkko, H., Rantanen, M., Paulin, L., Kotilainen, A., 2013. Sediment bacterial communities reflect the history of a sea basin. *PLoS One* 8, e54326.
- Major, C.O., Goldstein, S.L., Ryan, W.B., Lericolais, G., Piotrowski, A.M., Hajdas, I., 2006. The co-evolution of Black Sea level and composition through the last deglaciation and its paleoclimatic significance. *Quat. Sci. Rev.* 25, 2031–2047.
- Manske, A.K., Hensgen, U., Glaeser, J., Overmann, J., 2008. Subfossil 16S rRNA gene sequences of green sulfur bacteria in the Black Sea and their implications for past photic zone anoxia. *Appl. Environ. Microbiol.* 74, 624–632.
- Marcus, N.H., 1996. Ecological and evolutionary significance of resting eggs in marine

- copepods: past, present, and future studies. *Hydrobiologia* 320, 141–152.
- McConnell, J.R., Edwards, R., Kok, G.L., Flanner, M.G., Zender, C.S., Saltzman, E.S., Banta, J.R., Pasteris, D.R., Carter, M.M., Kahl, J.D., 2007. 20th-century industrial black carbon emissions altered arctic climate forcing. *Science* 317, 1381–1384.
- Miller, D.N., Bryant, J.E., Madsen, E.L., Ghiorse, W.C., 1999. Evaluation and optimization of DNA extraction and purification procedures for soil and sediment samples. *Appl. Environ. Microbiol.* 65, 4715–4724.
- Miteva, V., Sowers, T., Brechley, J., 2014. Penetration of fluorescent microspheres into the NEEM (North Eemian) Greenland ice core to assess the probability of microbial contamination. *Polar Biol.* 37, 47–59.
- Morard, R., Lejzerowicz, F., Darling, K.F., Lecroq-Bennet, B., Pedersen, M.W., Orlando, L., Pawlowski, J., Multiza, S., De Vargas, C., Kucera, M., 2017. Planktonic foraminifera-derived environmental DNA extract. *Biogeosciences* 14, 2741–2754.
- More, K.D., Orsi, W.D., Galy, V., Giosan, L., He, L., Grice, K., Coolen, M.J., 2018. A 43 kyr record of protist communities and their response to oxygen minimum zone variability in the Northeastern Arabian Sea. *Earth Planet. Sci. Lett.* 496, 248–256.
- Morono, Y., Inagaki, F., 2016. Chapter three - Analysis of Low-Biomass Microbial Communities in the Deep Biosphere. In: Sariaslani, S., Michael Gadd, G. (Eds.), *Advances in Applied Microbiology*. Academic Press, pp. 149–178.
- Negandhi, K., Laurion, I., Lovejoy, C., 2016. Temperature effects on net greenhouse gas production and bacterial communities in arctic thaw ponds. *FEMS Microbiol. Ecol.* 92, fiw117.
- Orcutt, B.N., Bergenthal, M., Freudenthal, T., Smith, D., Lilley, M.D., Schnieders, L., Green, S., Früh-Green, G.L., 2017. Contamination tracer testing with seabed drills: IODP Expedition 357. *Sci. Drill.* 23, 39–46.
- Orlando, L., Ginolhac, A., Zhang, G., Froese, D., Albrechtsen, A., Stiller, M., Schubert, M., Cappellini, E., Petersen, B., Moltke, I., Johnson, P.L., 2013. Recalibrating Equus evolution using the genome sequence of an early Middle Pleistocene horse. *Nature* 499, 74–78.
- Orsi, W.D., Coolen, M.J., Wuchter, C., He, L., More, K.D., Irigoien, X., Chust, G., Johnson, C., Hemingway, J.D., Lee, M., Galy, V., 2017. Climate oscillations reflected within the microbiome of Arabian Sea sediments. *Sci. Rep.* 7, 6040.
- Overballe-Petersen, S., Willerslev, E., 2014. Horizontal transfer of short and degraded DNA has evolutionary implications for microbes and eukaryotic sexual reproduction. *Bioessays* 36, 1005–1010.
- Pansu, J., Giguet-Covex, C., Ficetola, G.F., Gielly, L., Boyer, F., Zinger, L., Arnaud, J., Poulenard, Taberlet, P., Choler, P., 2015. Reconstructing long-term human impacts on plant communities: an ecological approach based on lake sediment DNA. *Mol. Ecol.* 24, 1485–1498.
- Pape, T., Hohnberg, H.J., Wunsch, D., Anders, E., Freudenthal, T., Huhn, K., Bohrmann, G., 2017. Design and deployment of autoclave pressure vessels for the portable deep-sea drill rig MeBo (Meeresboden-Bohrgerät). *Sci. Drill.* 23, 29–37.
- Parducci, L., Bennett, K.D., Ficetola, G.F., Alsos, I.G., Suyama, Y., Wood, J.R., Pedersen, M.W., 2017. Ancient plant DNA in lake sediments. *New Phytol.* 214 (3), 924–942.
- Parkes, R.J., Cragg, B.A., Bale, S.J., Getliff, J.M., Goodman, K., Rochelle, P.A., Fry, J.C., Weightman, A.J., Harvey, S.M., 1994. Deep bacterial biosphere in Pacific-Ocean sediments. *Nature* 371, 410–413.
- Pawlowska, J., Lejzerowicz, F., Esling, P., Szczuciński, W., Zajaczkowski, M., Pawlowski, J., 2014. Ancient DNA sheds new light on the Svalbard foraminiferal fossil record of the last millennium. *Geobiology* 12, 277–288.
- Pedersen, M.W., Overballe-Petersen, S., Ermini, L., Der Sarkissian, C., Haile, J., Hellstrom, M., Spens, J., Thomsen, P.F., Bohmann, K., Cappellini, E., Schnell, I.B., 2015. Ancient and modern environmental DNA. *Philos. Trans. R. Soc. B* 370, 20130383.
- Polz, M.F., Cavanaugh, C.M., 1998. Bias in template-to-product ratios in multitemplate PCR. *Appl. Environ. Microbiol.* 64, 3724–3730.
- Poté, J., Rosselli, W., Wigger, A., Wildi, W., 2007. Release and leaching of plant DNA in unsaturated soil column. *Ecotoxicol. Environ. Saf.* 68, 293–298.
- Randlett, M.E., Coolen, M.J.L., Stockhecke, M., Pickarski, N., Litt, T., Balkema, C., Kwiecien, O., Tomonaga, Y., Wehrli, B., Schubert, C.J., 2014. Alkenone distribution in Lake Van sediment over the last 270 ka: influence of temperature and haptophyte species composition. *Quat. Sci. Rev.* 10, 53–62.
- Rasmussen, M., Li, Y., Lindgreen, S., Pedersen, J.S., Albrechtsen, A., Moltke, I., Metspalu, M., Metspalu, E., Kivisild, T., Gupta, R., Bertalan, M., 2010. Ancient human genome sequence of an extinct Palaeo-Eskimo. *Nature* 463, 757.
- Rawlence, N.J., Lowe, D.J., Wood, J.R., Young, J.M., Churchman, G.J., Huang, Y.T., Cooper, A., 2014. Using palaeoenvironmental DNA to reconstruct past environments: progress and prospects. *J. Quat. Sci.* 29, 610–626.
- Rummel, J.D., Conley, C.A., 2017. Four fallacies and an oversight: searching for martian life. *Astrobiology* 17, 971–974.
- Salter, S.J., Cox, M.J., Turek, E.M., Calus, S.T., Cookson, W.O., Moffatt, M.F., Turner, P., Parkhill, J., Loman, N.J., Walker, A.W., 2014. Reagent and laboratory contamination can critically impact sequence-based microbiome analyses. *BMC Biol.* 12, 87.
- Schubert, M., Ermini, L., Sarkissian, C.D., Jónsson, H., Ginolhac, A., Schaefer, R., Martin, M.D., Fernández, R., Kircher, M., McCue, M., Willerslev, E., Orlando, L., 2014. Characterization of ancient and modern genomes by SNP detection and phylogenomic and metagenomic analysis using PALEOMIX. *Nat. Protoc.* 9, 1056–1082.
- Shapiro, B., Hofreiter, M., 2012. Preface. In: Shapiro, B., Hofreiter, M. (Eds.), *Ancient DNA, Methods and Protocols*. Springer New York, Dordrecht, Heidelberg, London, pp. v–vii.
- Sinninghe Damsté, J.S., Muyzer, G., Abbas, B., Rampen, S.W., Masse, G., Allard, W.G., Belt, S.T., Robert, J.M., Rowland, S.J., Moldovan, J.M., Barbanti, S.M., Fago, F.J., Denisevich, P., Dahl, J., Trindade, L.A.F., Schouten, S., 2004. The rise of the rhizosolenid diatoms. *Science* 304, 584–587.
- Slon, V., Hopfe, C., Weif, C.L., Mafessoni, F., De la Rasilla, M., Lalueza-Fox, C., Rosas, A., Soressi, M., Knul, M.V., Miller, R., Stewart, J.R., 2017. Neandertal and Denisovan DNA from Pleistocene sediments. *Science* 356, 605–608.
- Smith, D.C., Spivack, A.J., Fisk, M.R., Haveman, S.A., Staudigel, H., 2000. Tracer-based estimates of drilling-induced microbial contamination of deep sea crust. *Geomicrobiol. J.* 17, 207–219.
- Stat, M., Huggett, M.J., Bernasconi, R., DiBattista, J.D., Berry, T.E., Newman, S.J., Harvey, E.S., Bunce, M., 2017. Ecosystem biomonitoring with eDNA: metabarcoding across the tree of life in a tropical marine environment. *Sci. Rep.* 7, 12240.
- Sunagawa, S., Coelho, L.P., Chaffron, S., Kultima, J.R., Labadie, K., Salazar, G., Djahanschiri, B., Zeller, G., Mende, D.R., Alberti, A., Cornejo-Castillo, F.M., 2015. Structure and function of the global ocean microbiome. *Science* 348, 1261359.
- Szczuciński, W., Pawlowska, J., Lejzerowicz, F., Nishimura, Y., Kokociński, M., Majewski, W., Nakamura, Y., Pawlowski, J., 2016. Ancient sedimentary DNA reveals past tsunami deposits. *Mar. Geol.* 381, 29–33.
- Taberlet, P., Coissac, E., Hajibabaei, M., Rieseberg, L.H., 2012a. Environmental DNA. *Mol. Ecol.* 21, 1789–1793.
- Taberlet, P., Prud'homme, S.M., Campione, E., Roy, J., Miquel, C., Shehzad, W., Gielly, L., Rioux, D., Choler, P., Clément, J.C., 2012b. Soil sampling and isolation of extracellular DNA from large amount of starting material suitable for metabarcoding studies. *Mol. Ecol.* 21, 1816–1820.
- Tada, R., Murray, R., Alvarez Zarikian, C., Anderson Jr., W., Bassetti, M., Brace, B., Clemens, S., da Costa Gurgel, M., Dickens, G., Dunlea, A., 2013. Asian Monsoon: onset and evolution of millennial-scale variability of Asian monsoon and its possible relation with Himalaya and Tibetan Plateau uplift. *IODP Scientific Prospect.* 346, 1–111.
- Torres, M.E., Cox, T.L., Hong, W.-L., McManus, J., Sample, J.C., Desgrigneve, C., Gan, H.M., Moreau, J.M., 2015. Crustal fluid and ash alteration impacts on the biosphere of Shikoku Basin sediments, Nankai Trough, Japan. *Geobiology* 13, 562–580.
- Torti, A., Lever, M.A., Jørgensen, B.B., 2015. Origin, dynamics, and implications of extracellular DNA pools in marine sediments. *Mar. Genomics* 24, 185–196.
- Van der Meer, M.T.J., Sangiorgi, F., Baas, M., Brinkhuis, H., Sinninghe Damsté, J.S., Schouten, S., 2008. Molecular isotopic and dinoflagellate evidence for late Holocene freshening of the Black Sea. *Earth Planet. Sci. Lett.* 267, 426–434. Van der Meer, M.T.J., Sangiorgi, F., Baas, M., Brinkhuis, H., Sinninghe Damsté, J.S., and Schouten, S.
- Volkman, J.K., Barrett, S.M., Blackburn, S.I., Mansour, M.P., Sikes, E.L., Gelin, F., 1998. Microalgal biomarkers: a review of recent research developments. *Org. Geochem.* 29, 1163–1179.
- Wagner, A., Blackstone, N., Cartwright, P., Dick, M., Misof, B., Snow, P., Wagner, G.P., Bartels, J., Murtha, M., Pendleton, J., 1994. Surveys of gene families using polymerase chain reaction: PCR selection and PCR drift. *Syst. Biol.* 43, 250–261.
- Webster, G., Newberry, C.J., Fry, J.C., Weightman, A.J., 2003. Assessment of bacterial community structure in the deep sub-seafloor biosphere by 16S rDNA-based techniques: a cautionary tale. *J. Microbiol. Methods* 55, 155–164.
- Weyrich, L.S., Dobney, K., Cooper, A., 2015. Ancient DNA analysis of dental calculus. *J. Hum. Evol.* 79, 119–124.
- Weyrich, L.S., Duchene, S., Soubrier, J., Arriola, L., Llamas, B., Breen, J., Morris, A.G., Alt, K.W., Caramelli, D., Dresely, V., 2017. Neandertal behaviour, diet, and disease inferred from ancient DNA in dental calculus. *Nature* 544, 357.
- Willerslev, E., Cooper, A., 2005. Ancient DNA. *Proc. R. Soc. Lond. B Biol. Sci.* 272, 3–16.
- Willerslev, E., Hansen, A.J., Binladen, J., Brand, T.B., Gilbert, M.T.P., Shapiro, B., Bunce, M., Wiuf, C., Gilichinsky, D.A., Cooper, A., 2003. Diverse plant and animal genetic records from Holocene and Pleistocene sediments. *Science* 300, 791–795.
- Willerslev, E., Hansen, A.J., Poinar, H.N., 2004. Isolation of nucleic acids and cultures from fossil ice and permafrost. *Trends Ecol. Evol.* 19, 141–147.
- Willerslev, E., Davison, J., Moora, M., Zobel, M., Coissac, E., Edwards, M.E., Lorenzen, E.D., Vestergård, M., Gussarova, G., Haile, J., Craine, J., Gielly, L., Boessenkool, S., Epp, L.S., Pearnan, P.B., Cheddadi, R., Murray, D., Brathen, K.A., Yoccoz, N., Binney, H., Craud, C., Wincker, P., Goslar, T., Alsos, I.G., Bellemain, E., Brysting, A.K., Elven, R., Sonstebo, J.H., Murton, J., Sher, A., Rasmussen, M., Ronn, R., Mourier, T., Cooper, A., Austin, J., Moller, P., Froese, D., Zazula, G., Pompanon, F., Rioux, D., Niderkorn, V., Tikhonov, A., Savvinov, G., Roberts, R.G., MacPhee, R.D., Gilbert, M.T., Kjaer, K.H., Orlando, L., Brochmann, C., Taberlet, P., 2014. Fifty thousand years of Arctic vegetation and megafaunal diet. *Nature* 506, 47–51.
- Winter, D., Spjurneskog, C., Harwood, D., 2010. Early to mid-Pliocene environmentally constrained diatom assemblages from the AND-1B drillcore. *McMurdo Sound, Antarctica. Stratigraphy* 7, 207–227.
- Young, J.M., Weyrich, L.S., Clarke, L.J., Cooper, A., 2015. Residual soil DNA extraction increases the discriminatory power between samples. *Forensic Sci. Med. Pathol.* 11, 268–272.
- Zhou, J., Bruns, M.A., Tiedje, J.M., 1996. DNA recovery from soils of diverse composition. *Appl. Environ. Microbiol.* 62, 316–322.
- Zinger, L., Amaral-Zettler, L.A., Fuhrman, J.A., Horner-Devine, M.C., Huse, S.M., Welch, D.B.M., Martiny, J.B., Sogin, M., Boetius, A., Ramette, A., 2011. Global patterns of bacterial beta-diversity in seafloor and seawater ecosystems. *PLoS One* 6, e24570.